

Effects of the changing geomagnetic field on the atmosphere

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Antarctic Survey**

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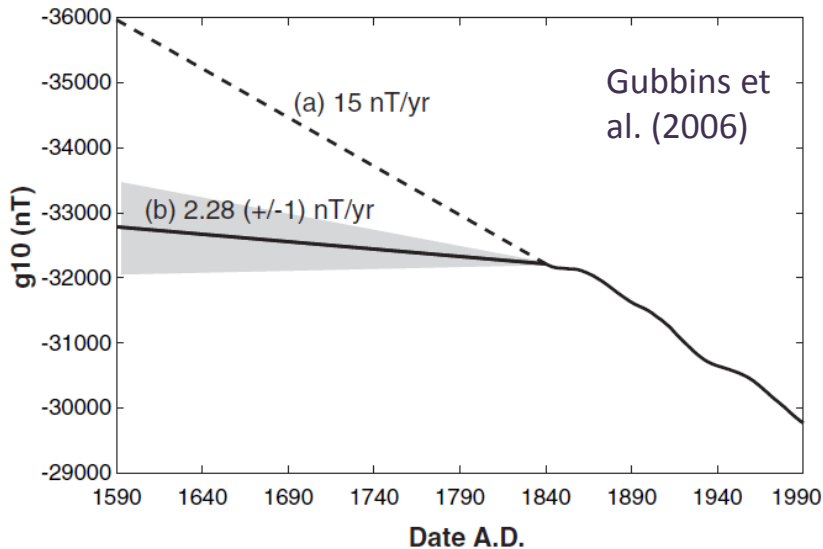
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POTSDAM

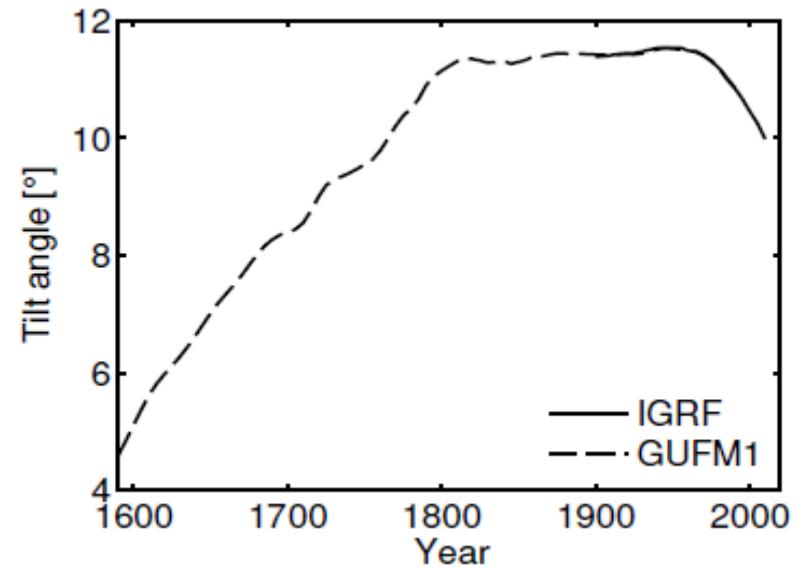
Outline

■ Magnetosphere and upper atmosphere

■ Variations in dipole field *strength*



■ Variations in dipole field *orientation*

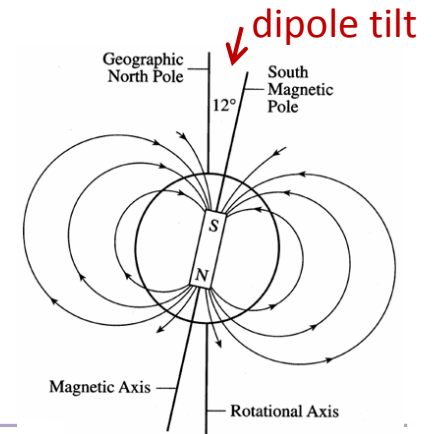


■ Realistic magnetic field changes 1908-2008

■ Contrasted with effects of CO₂ increase

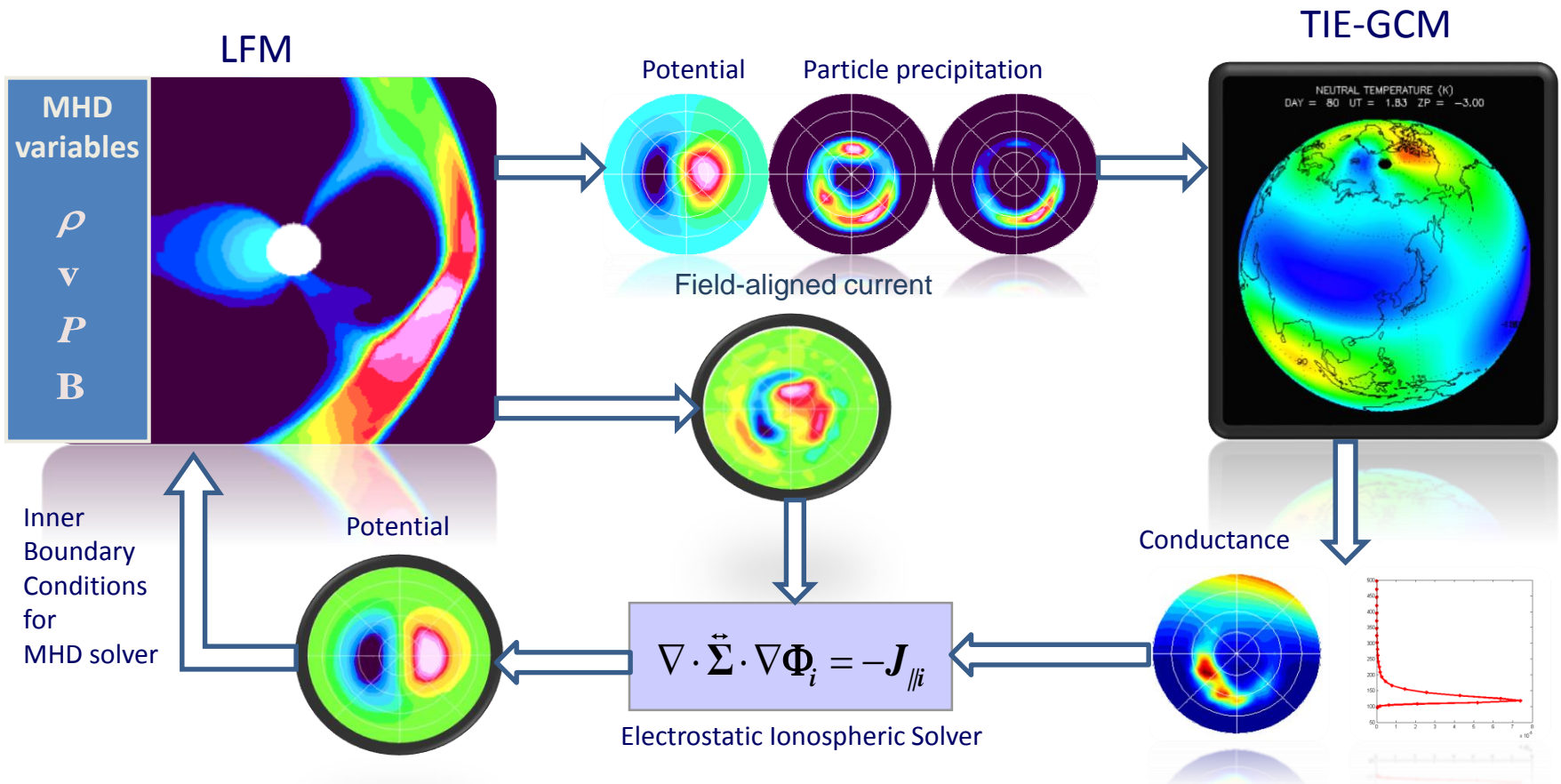
■ Lower and middle atmosphere

■ Realistic magnetic field changes 1900-2000



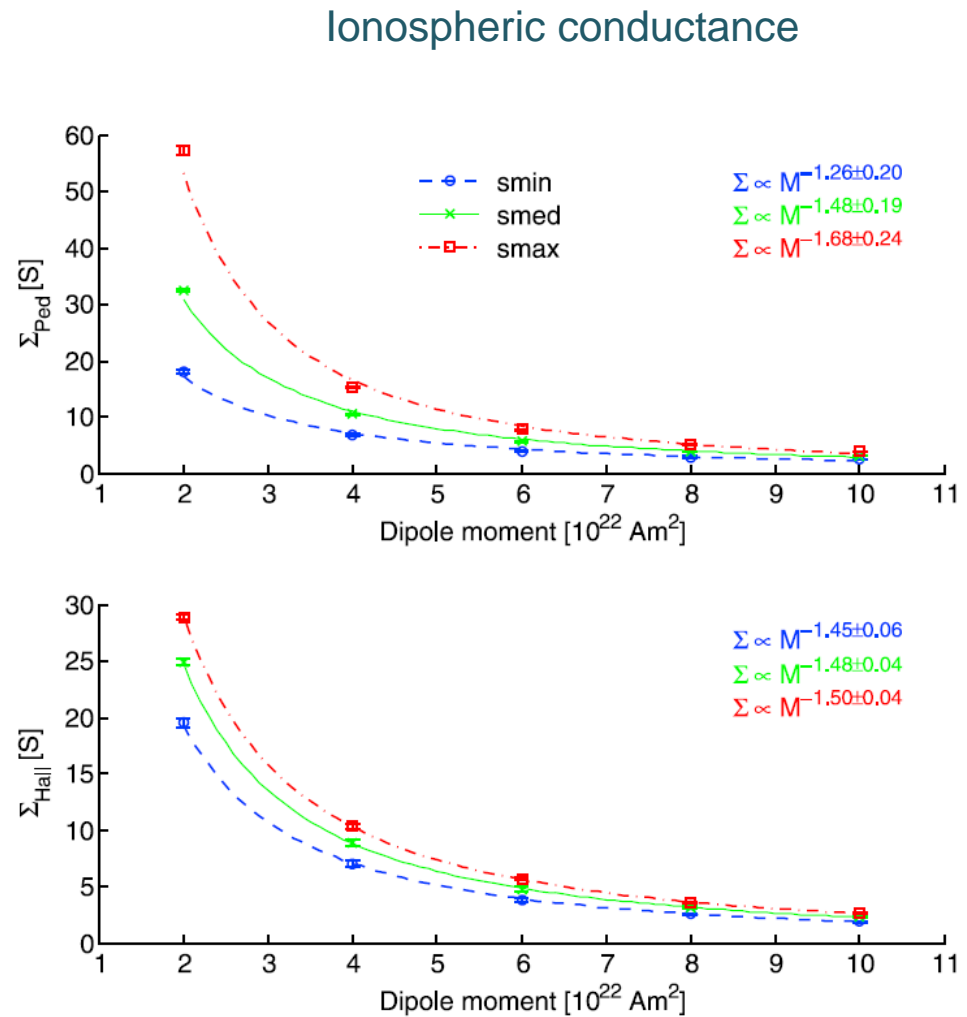
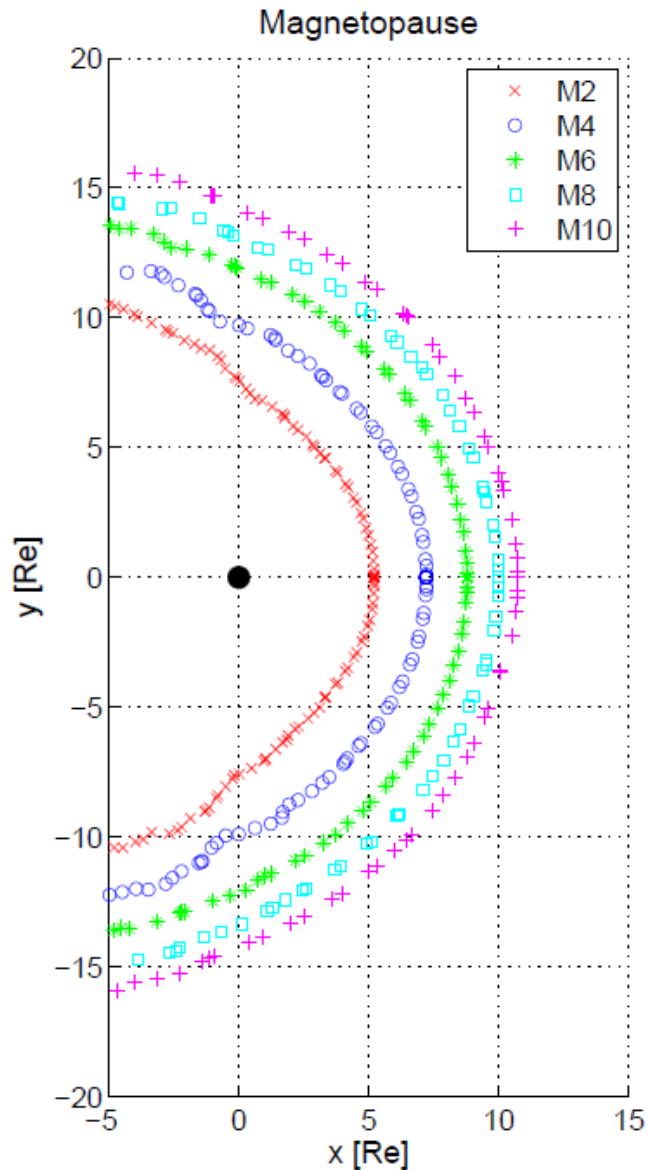
Coupled Magnetosphere-Ionosphere-Thermosphere model

- CMIT = LFM + TIE-GCM
- LFM = Lyon-Fedder-Mobarry MHD code (magnetosphere model)
- TIE-GCM = Thermosphere-Ionosphere-Electrodynamics General Circulation Model



Graphics courtesy of Binzheng Zhang

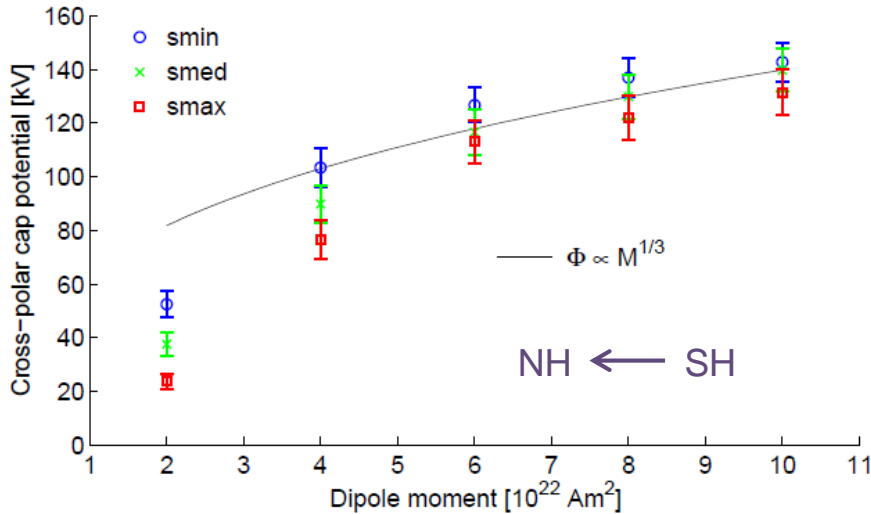
Influence of magnetic field strength



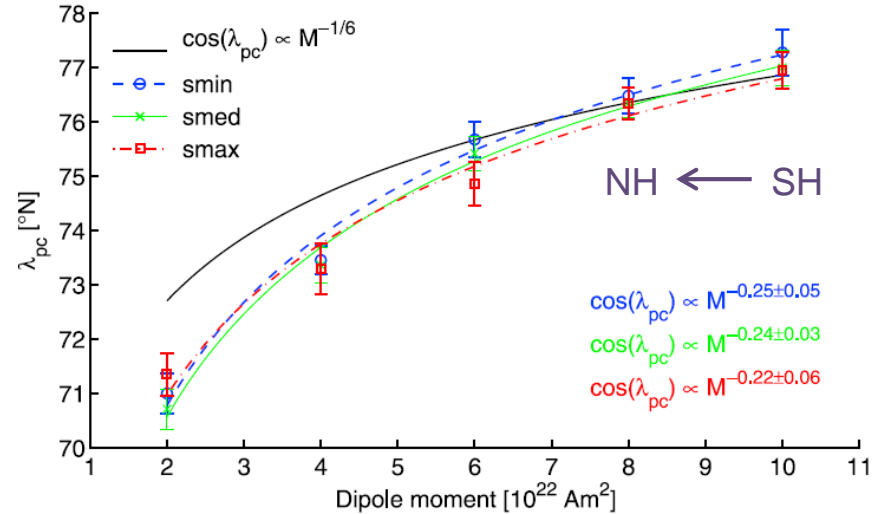
Crossen, Richmond and Wiltberger, JGR, 2012

Influence of magnetic field strength

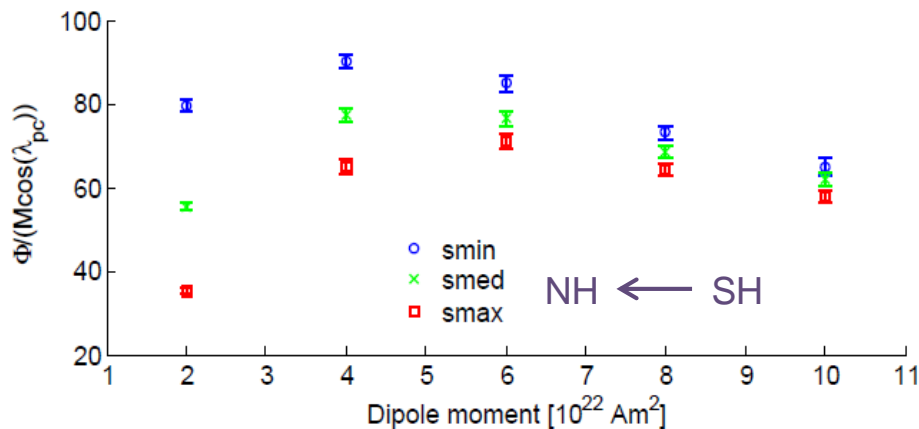
Cross-polar cap potential



Latitude of polar cap boundary



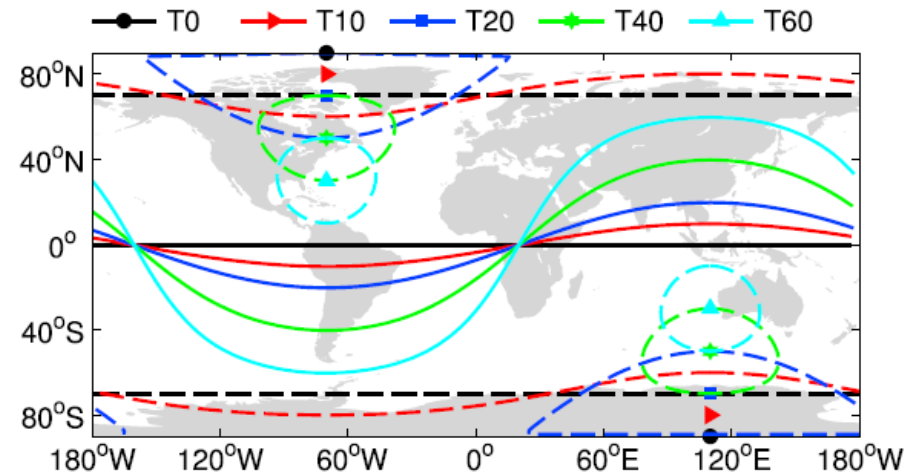
Measure of ExB drift magnitude (E/B)



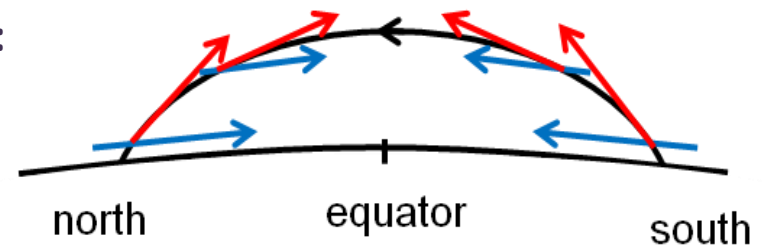
- For lower magnetic field strength
 - Cross-polar cap potential ↓
 - Polar cap size ↑
 - ExB drift magnitude ↑
- ExB drifts and high-latitude neutral winds NH>SH

Why the dipole orientation matters for the ionosphere

- Geographic locations of important features, e.g., magnetic equator, magnetic poles, auroral oval
- Orientation of the magnetic poles w.r.t. the Sun and solar wind
 - Solar illumination of polar cap
 - Solar wind-magnetosphere-ionosphere coupling efficiency (see Cnossen, Wiltberger and Ouellette, JGR, 2012)
 - Strength of polar ionospheric flows and Joule heating
- Plasma transport depends on orientation of **B**:
 - \mathbf{ExB} drifts
 - Plasma transport by neutral winds
 - Diffusion along the magnetic field



Cnossen and Richmond
(JGR, 2012)

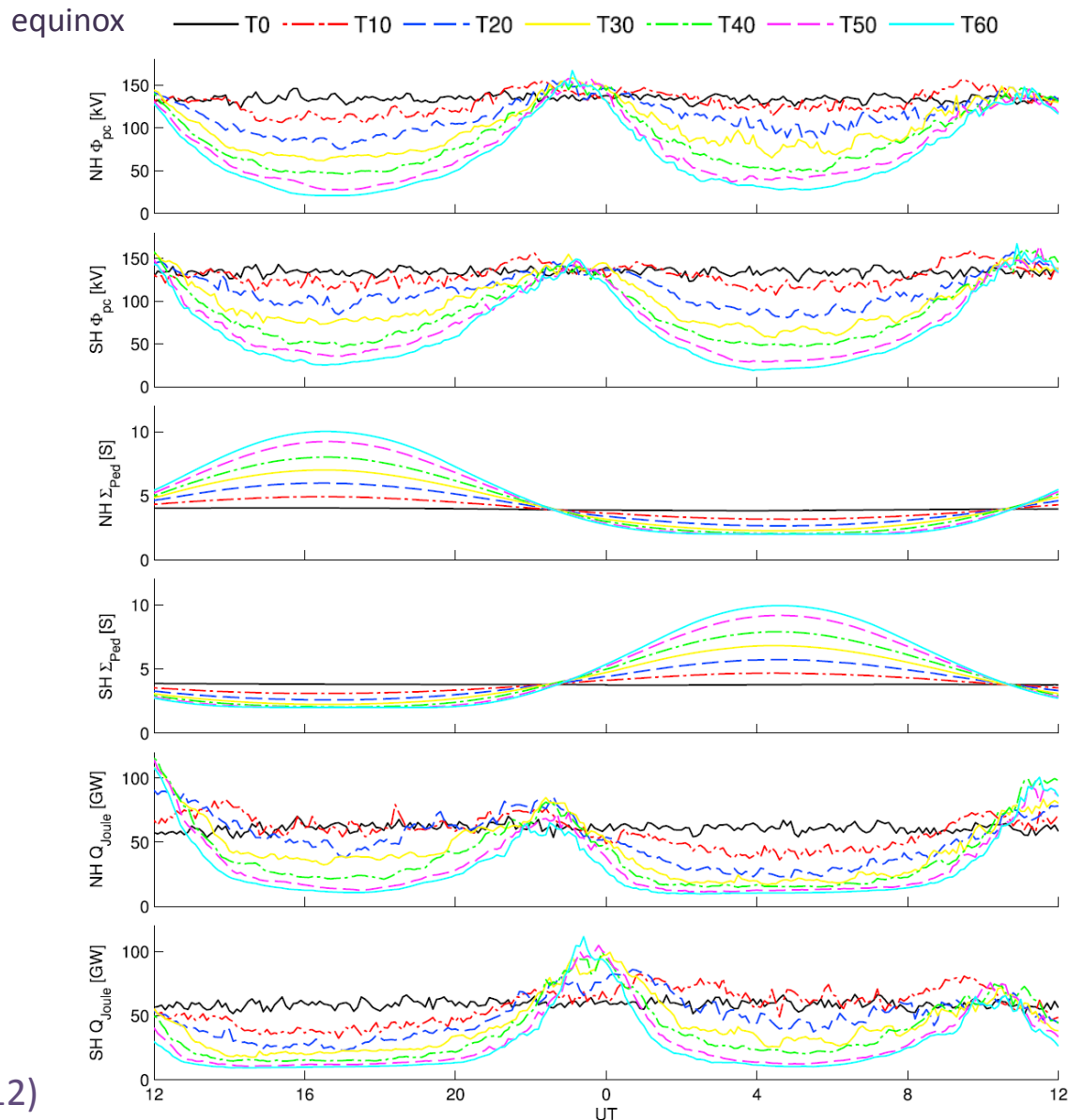


Cnossen (InTech Open, 2012)

Diurnal variation in CPCP, conductivity and Joule heating

Equinox:

- Larger diurnal variation for larger dipole tilt
- Lower mean CPCP and Joule heating for larger dipole tilt



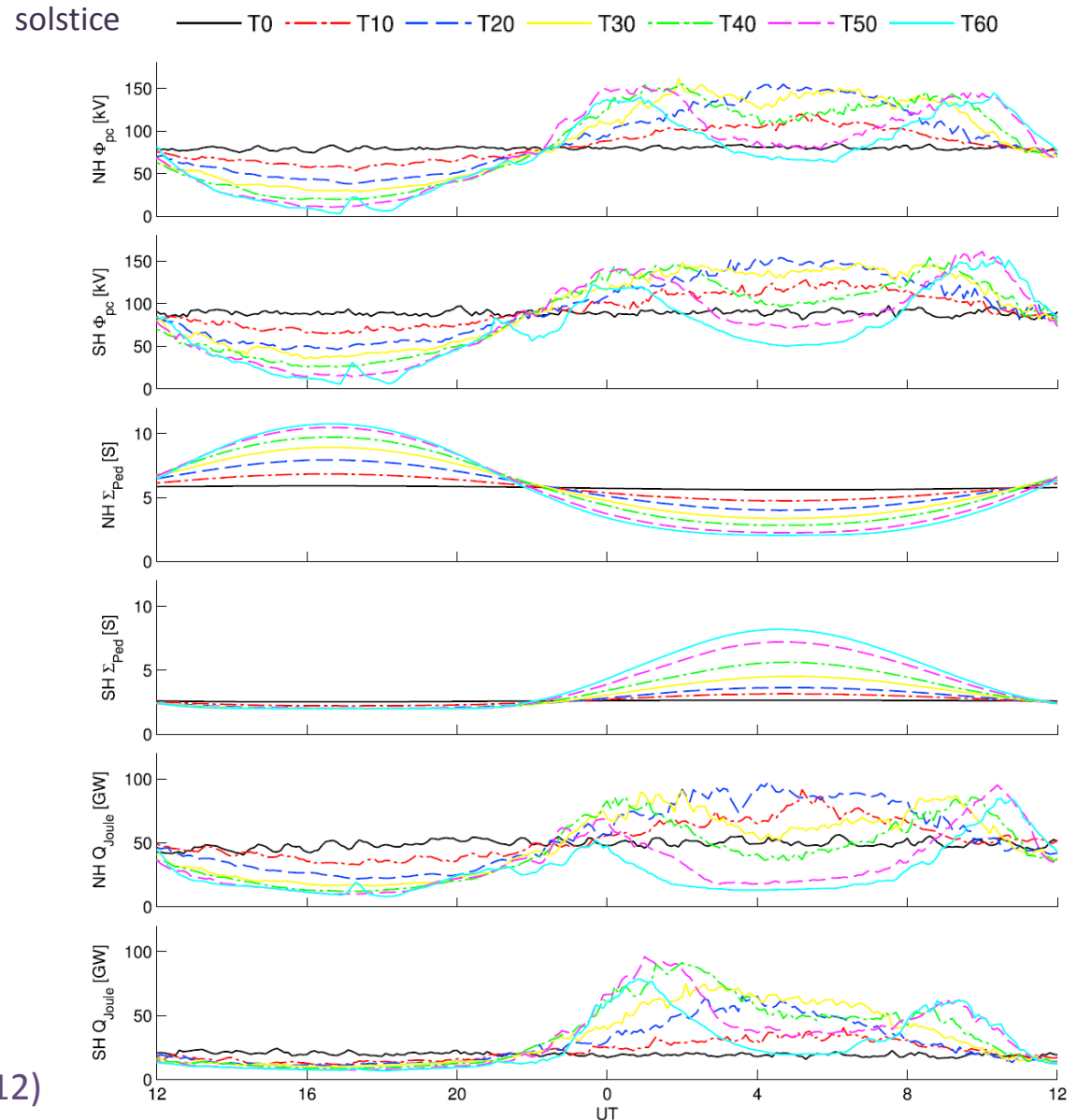
Diurnal variation in CPCP, conductivity and Joule heating

Equinox:

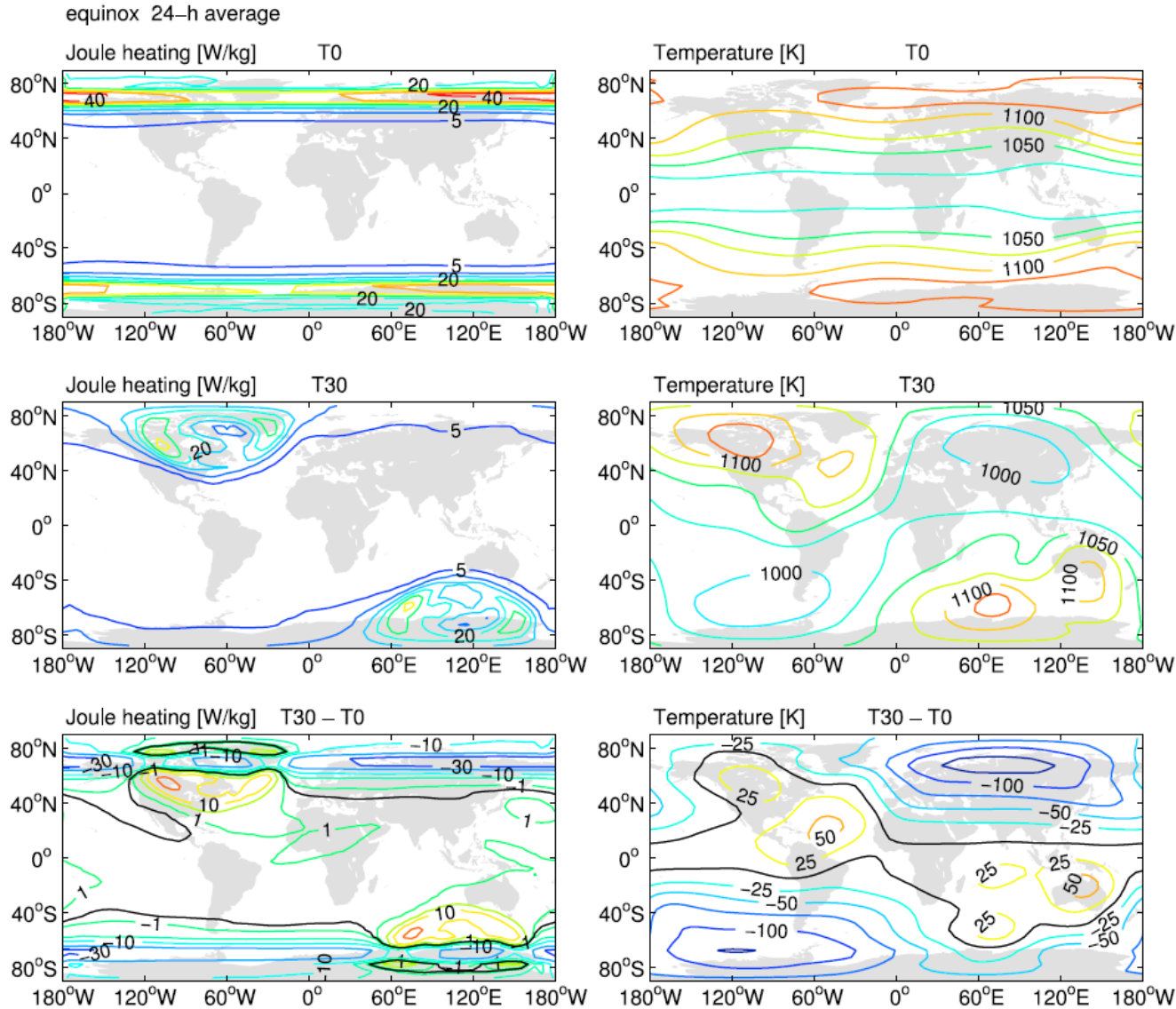
- Larger diurnal variation for larger dipole tilt
- Lower mean CPCP and Joule heating for larger dipole tilt

Solstice:

- More complicated diurnal variation pattern for large tilts due to multiple $\mu = 0$ crossings
- μ = angle between dipole axis and GSM z-axis



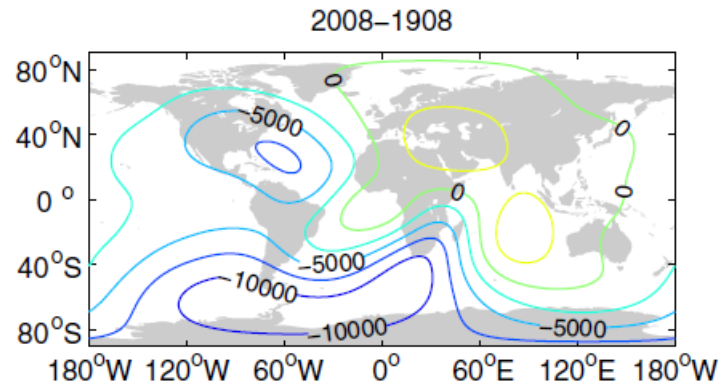
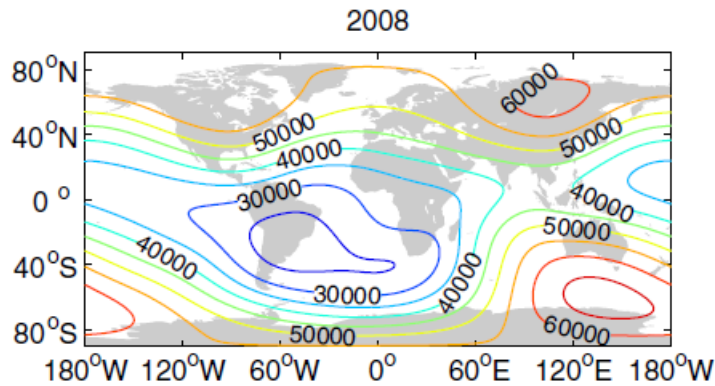
T0 vs. T30: Joule heating and temperature



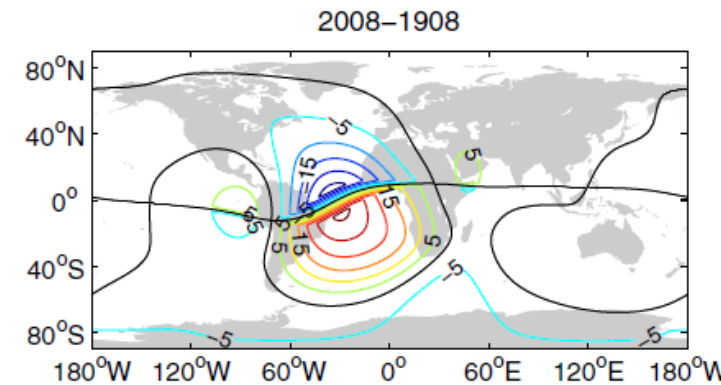
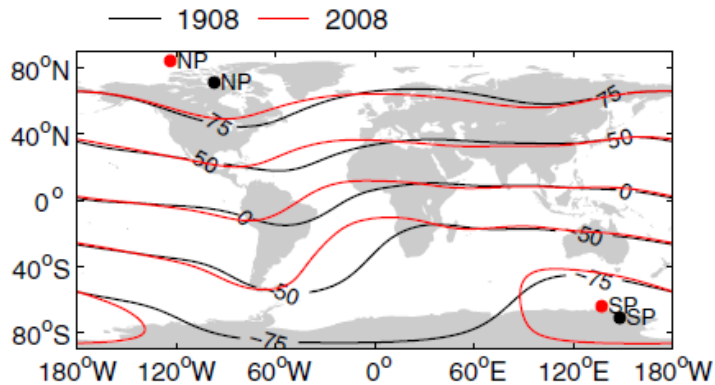
- T30 gives less Joule heating than T0
- Changes in geographic distribution of Joule heating
- Changes in temperature structure more or less follow

Crossen and Richmond
(JGR, 2012)

Realistic magnetic field changes 1908-2008



Magnetic field strength in 2008 and difference with 1908 (nT)



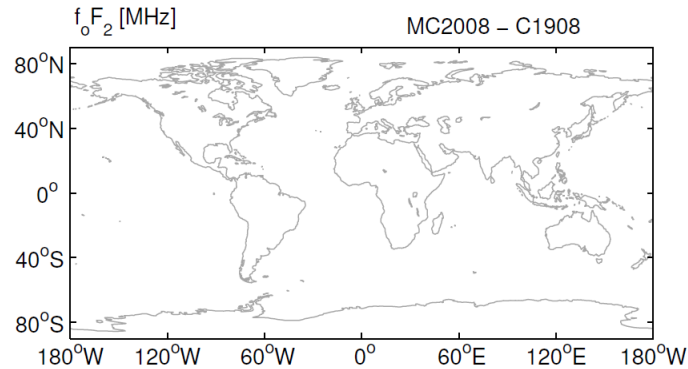
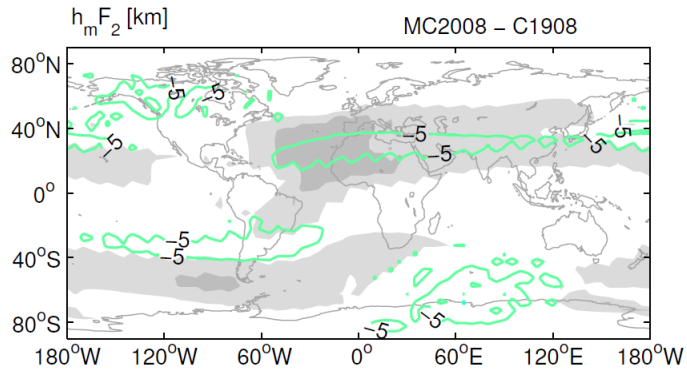
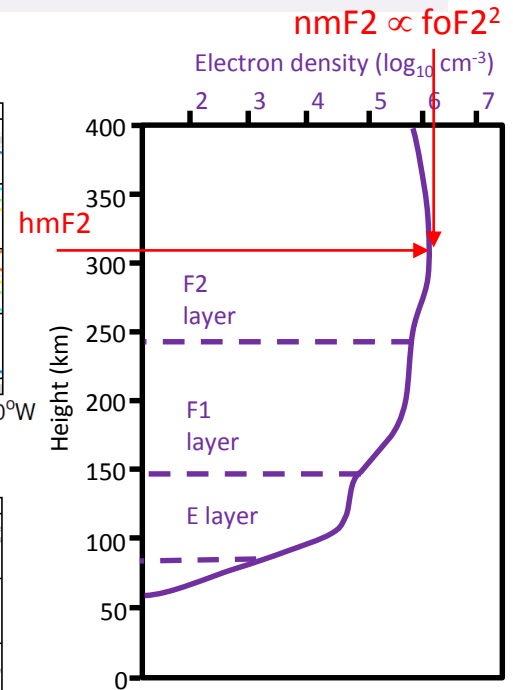
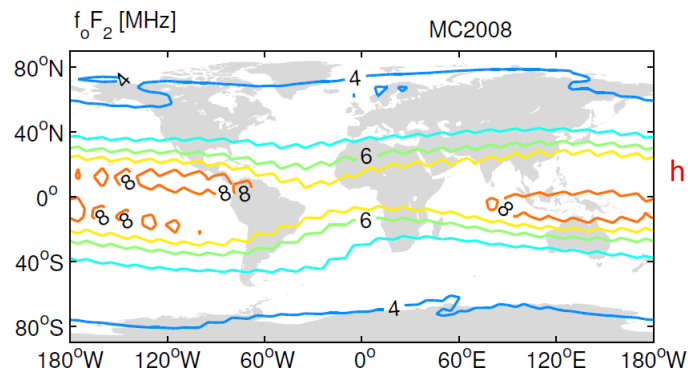
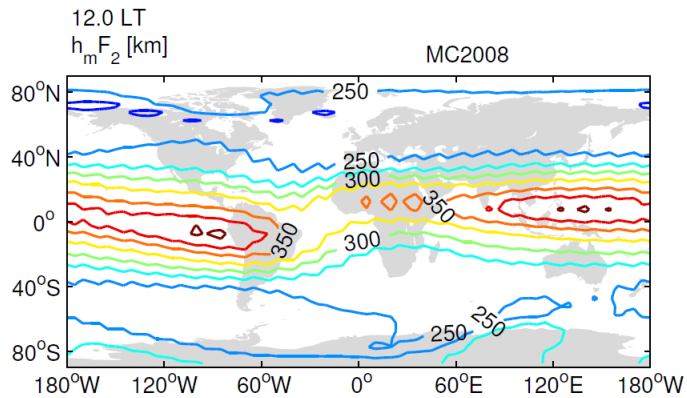
Inclination angle contours of 1908 and 2008 and differences in inclination (°)

- Expansion and intensification of the South Atlantic Anomaly region of low magnetic field strength
- Northward and westward movement of magnetic field structures
- Strongest inclination angle changes in Atlantic region (~100°W-50°E; ~40°S-40°N)

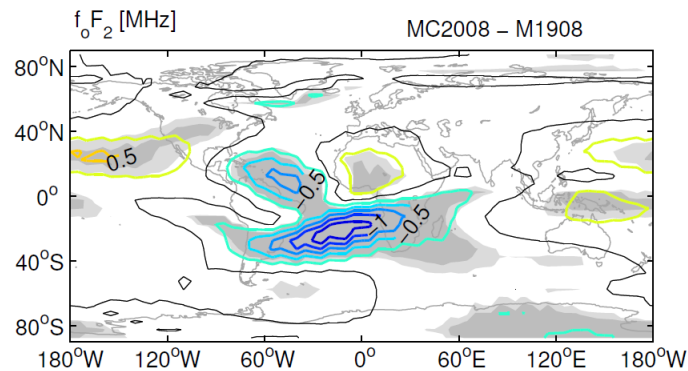
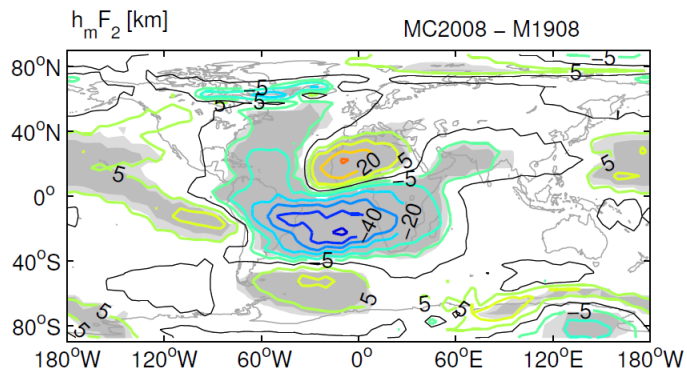
IGRF =
International
Geomagnetic
Reference Field

Cnossen and
Richmond (JGR, 2013)

Magnetic field vs. CO₂ effects: h_mF₂ and f_oF₂



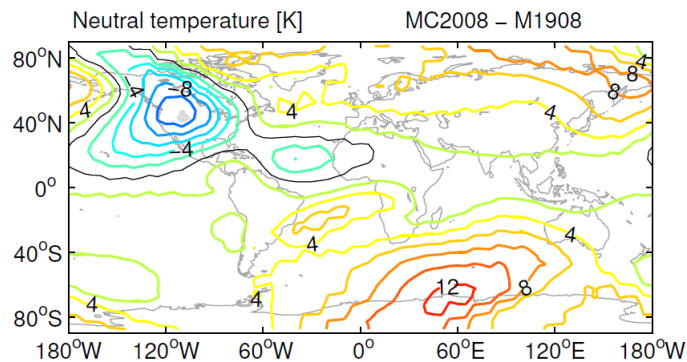
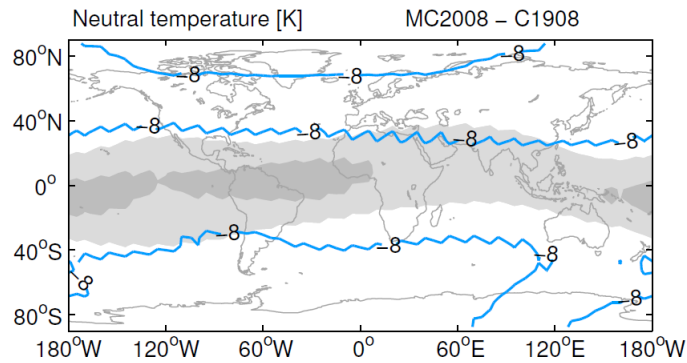
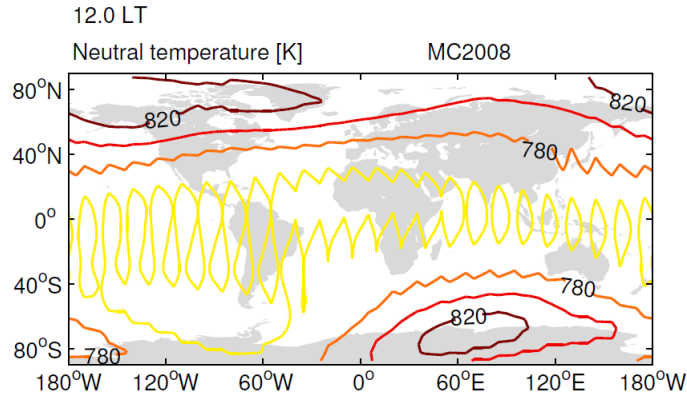
← CO₂ effect



← Magnetic field effect

Crossen (JSWSC, 2014)

Magnetic field vs. CO₂ effects: temperature @ 300 km

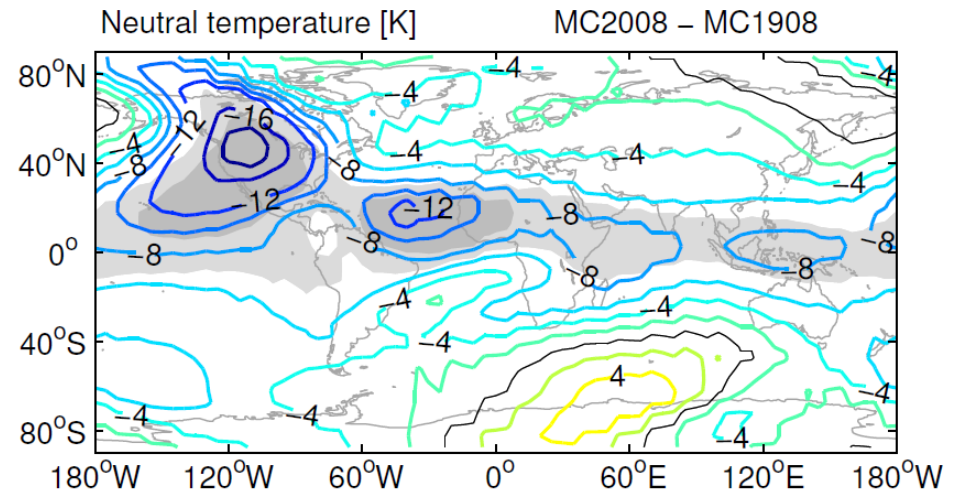


CO₂ effect

Magnetic field effect

- Increase in CO₂ concentration is more important
- But changes in the magnetic field do contribute

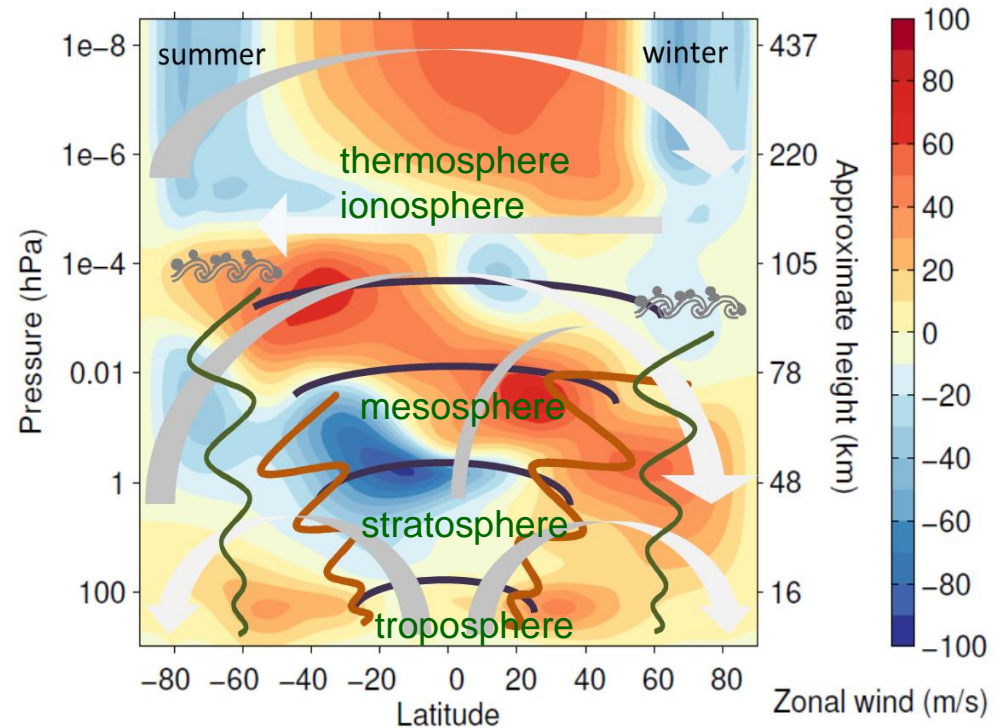
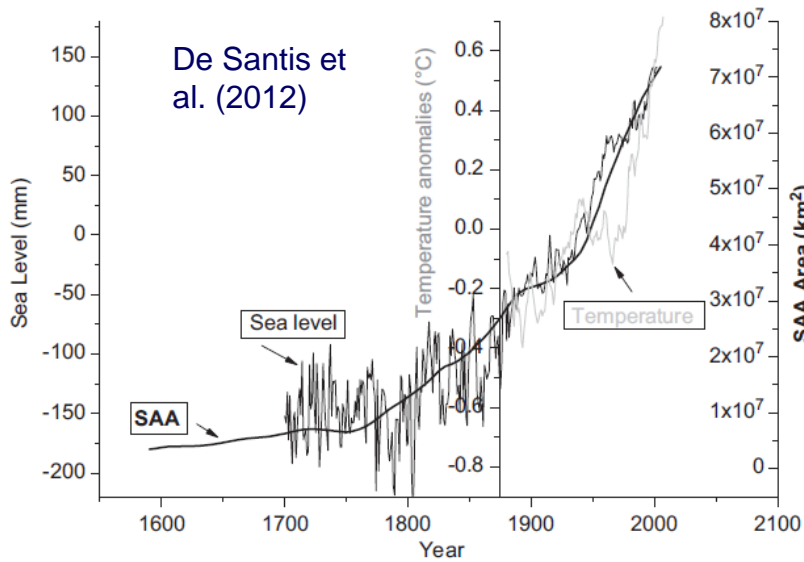
Combined CO₂ and magnetic field effect



Crossen (JSWSC, 2014)

Whole atmosphere response to magnetic field changes

- Some observational evidence for a magnetic field influence on tropospheric climate



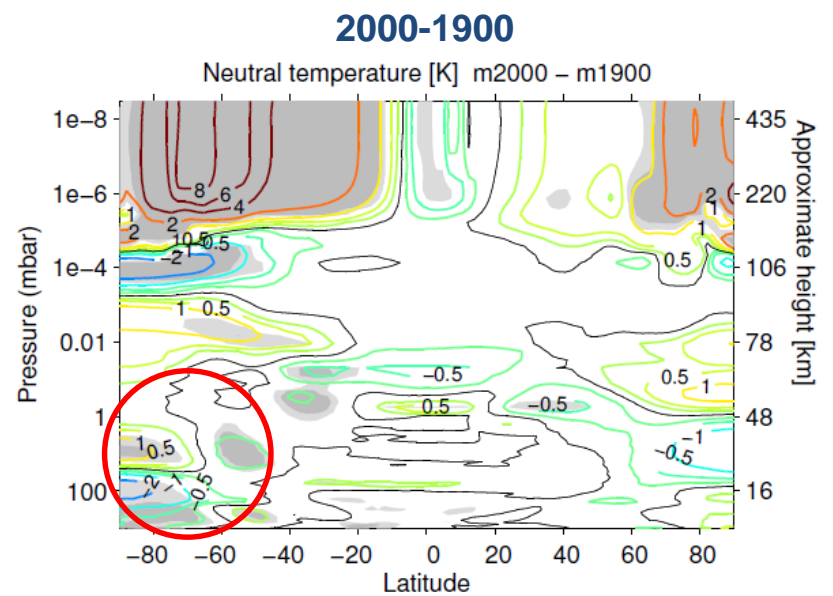
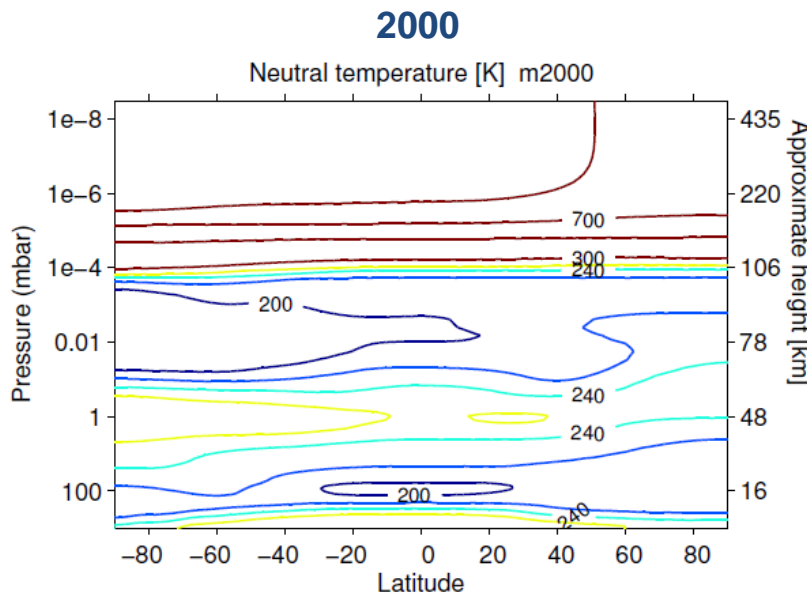
- Mechanism unclear – *controversial*
- **Whole atmosphere dynamical coupling?**
 - Upwardly propagating atmospheric waves induce a residual circulation that extends downwards when they break

Lower/middle atmosphere response (DJF)

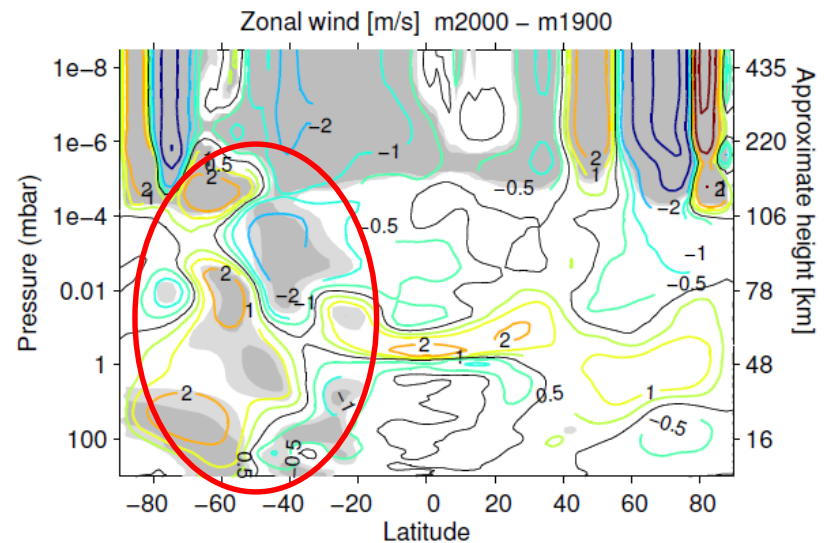
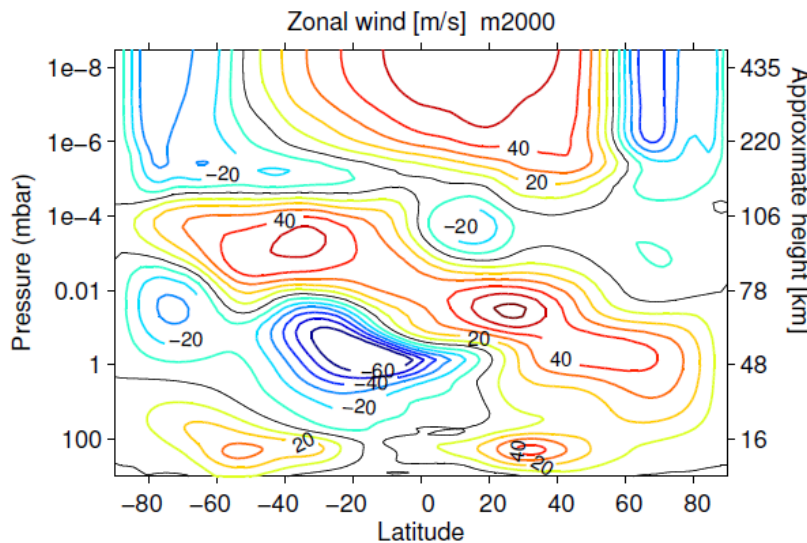
DJF

Zonal
mean

Temp.



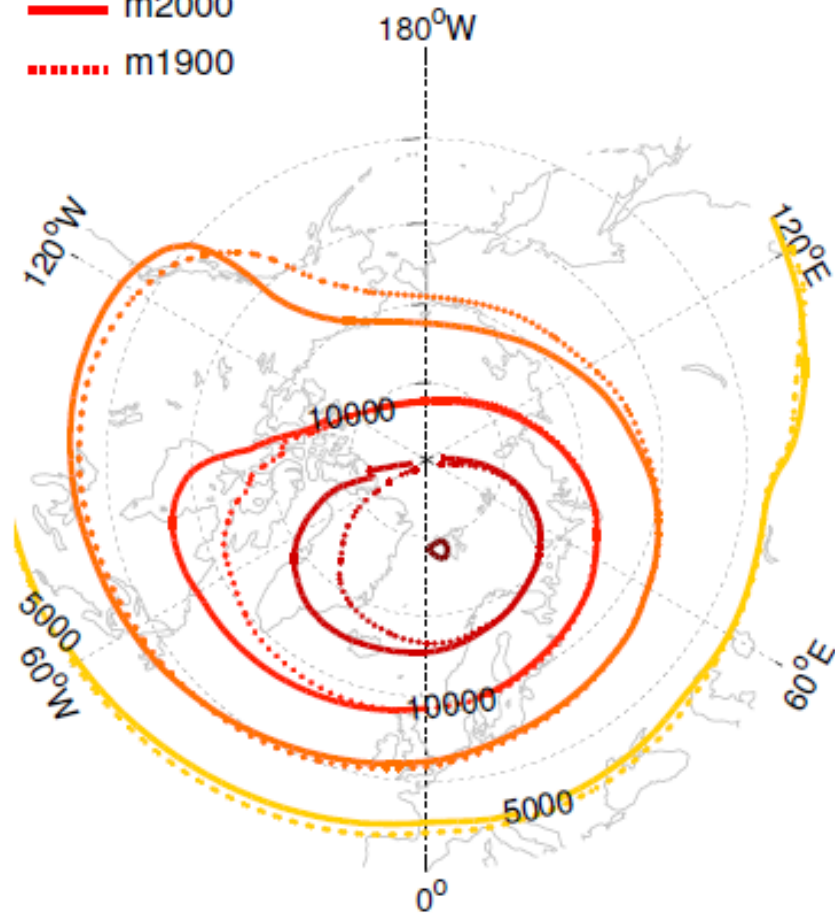
Zonal
wind



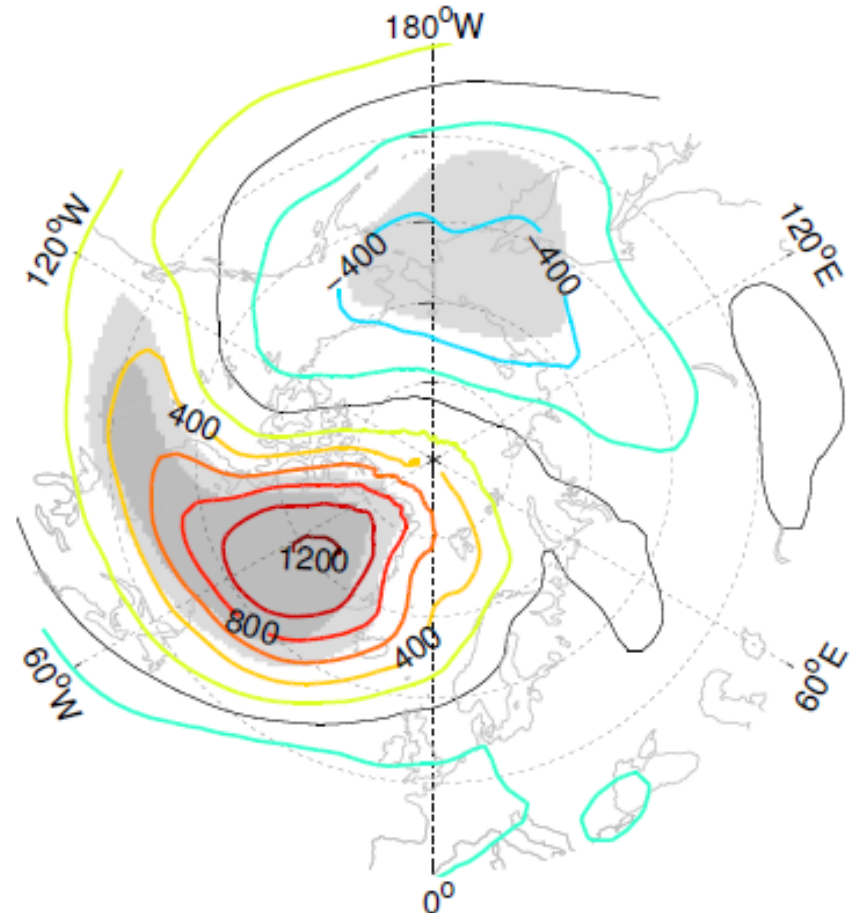
Movement/distortion of NH polar vortex

NH Potential vorticity (PV) [pvu] DJF 0.96 hPa

— m2000
- - - m1900



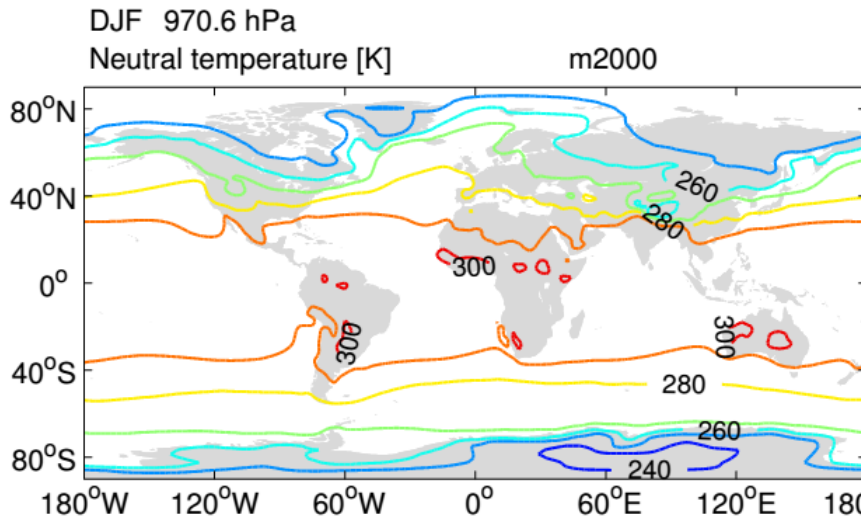
m2000-m1900



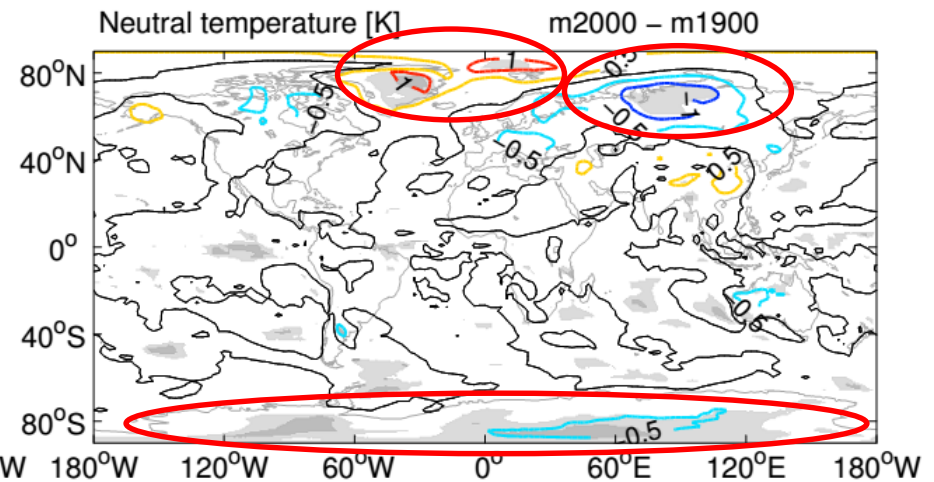
Crossen , Liu, and Lu (2016, in review)

Surface temperature

2000



2000-1900



- NH: significant warming over Greenland and cooling over Siberia (± 1 K)
 - Resembles negative phase of NAO
 - Inconsistent with changes in stratospheric vortex
- SH: significant high-latitude cooling (-0.5 K)

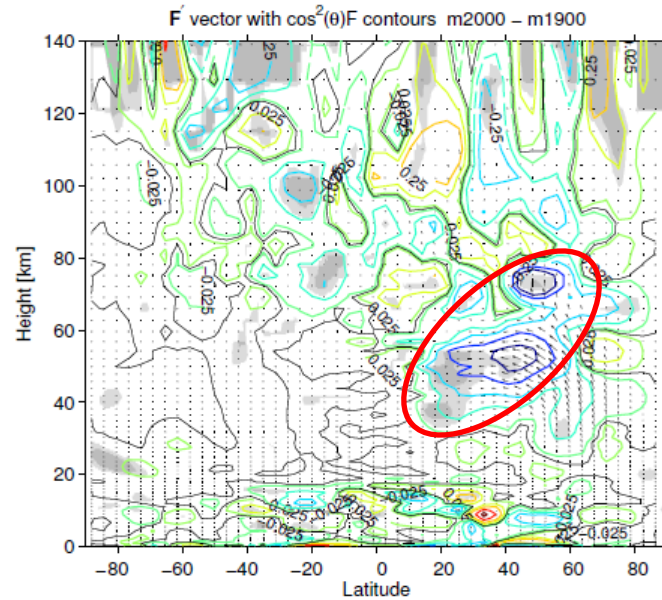
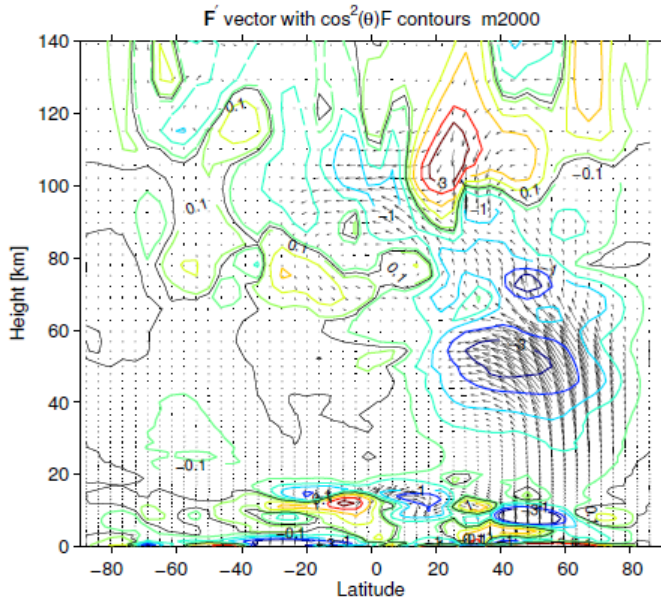
Summary and conclusions

- Changes in the Earth's main magnetic field cause significant long-term change in the upper atmosphere
 - Both field strength and orientation are important
 - Mechanisms mostly understood
 - Magnetic field changes are as important for upper atmosphere climate as the increase in CO₂ concentration!
 - Upper atmosphere changes can affect the atmosphere below via changes in wave forcing, which exert a downward influence
 - Dependent on season (strongest for DJF)
 - Dependent on longitude (in NH)
 - Vertical coupling mechanism is sensitive to (inaccuracies in) simulated background climatology – model improvements needed
-

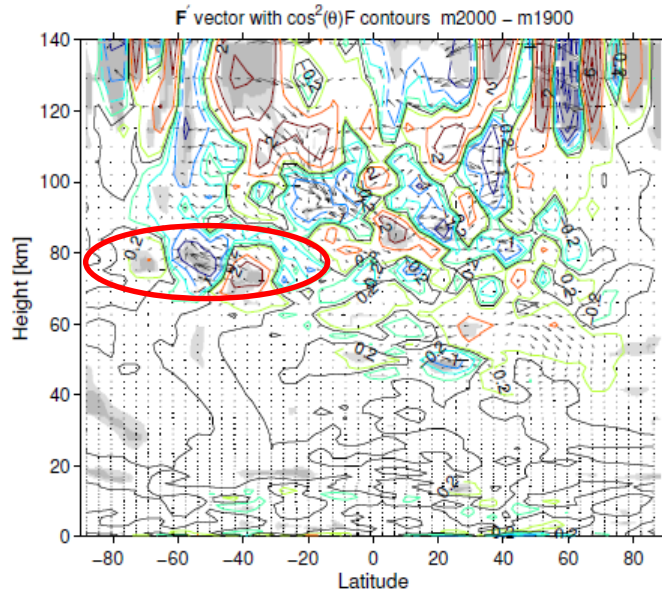
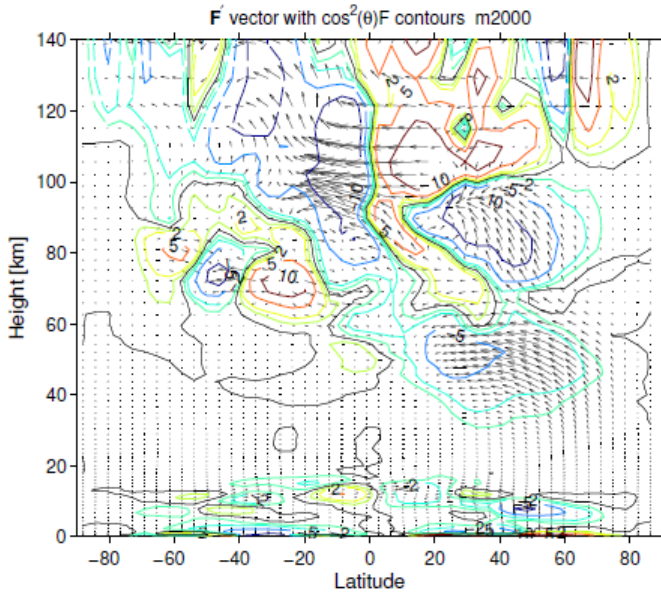
Changes in resolved wave forcing (EP flux)

DJF

from DJF
average data
(stationary
waves)



from 5-
daily data
(transient
waves)



Probably influence of quasi-two-day wave (QTDW)

