

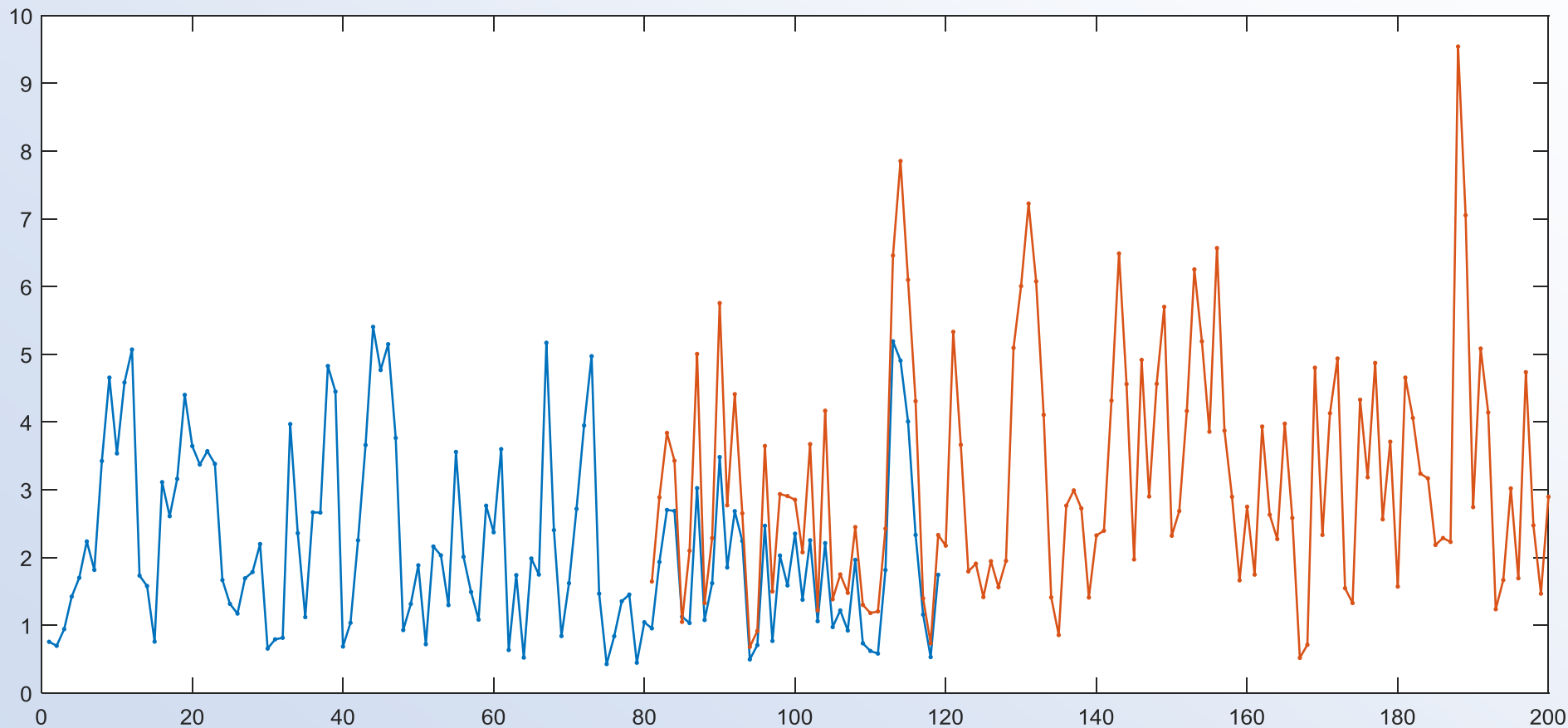


Long-term homogeneity and intercalibration of datasets

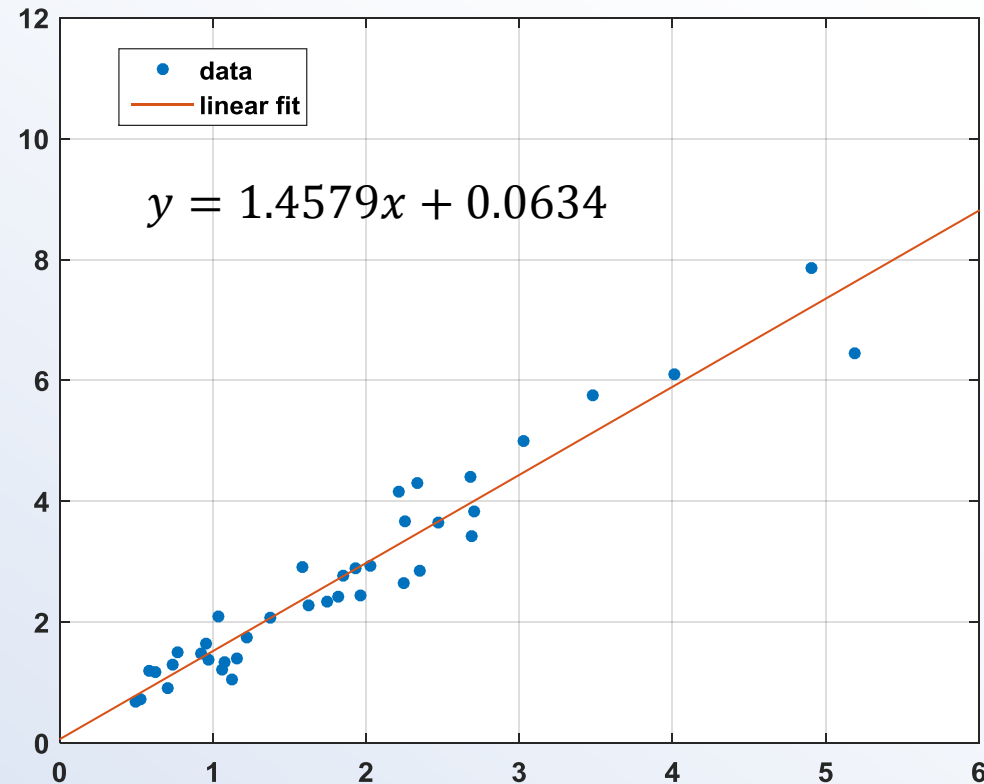
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- Why do we need to homogenize and intercalibrate?
- Differences/changes in instrumentation, location of measurements, method of measurement etc.
- Basic methods of forming homogeneous data and inter-calibrating between datasets
- **Homogenization** = making sure your data describes the same thing in same way at all times
- **Intercalibration** = making sure two (or more) data series describe the same thing in same way at all times
- Foundation: understand your data and measurements!
- Examples:
 - Geomagnetic indices
 - Particle measurements from satellites

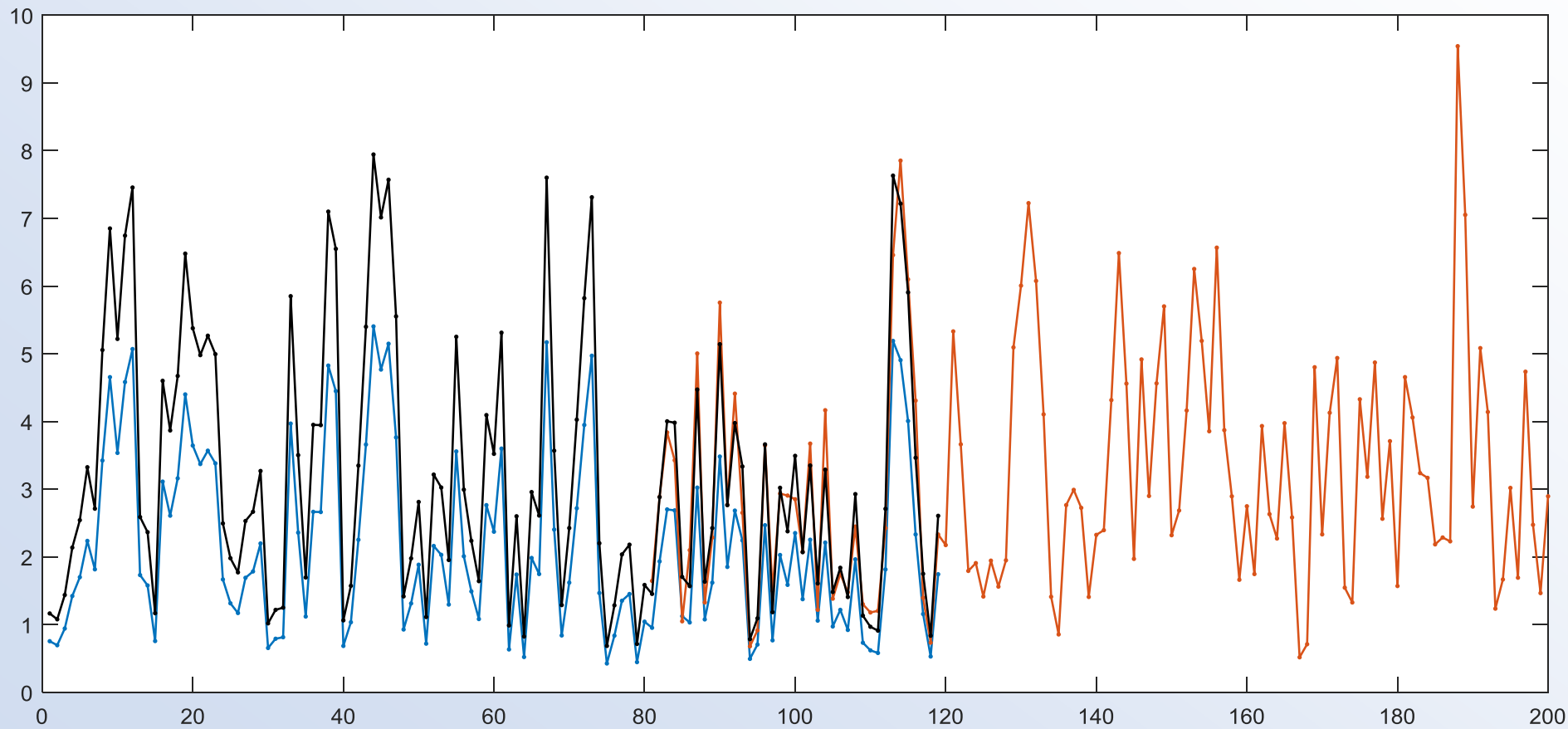
- Blue series ends and red series begins
- Overlapping data differs → two series systematically different → Compare overlapping part



- First make sure data are comparable!
- We assume here that individual series are homogeneous during their individual periodss
- Comparison by scatterplot → fit a suitable curve (e.g., a line)

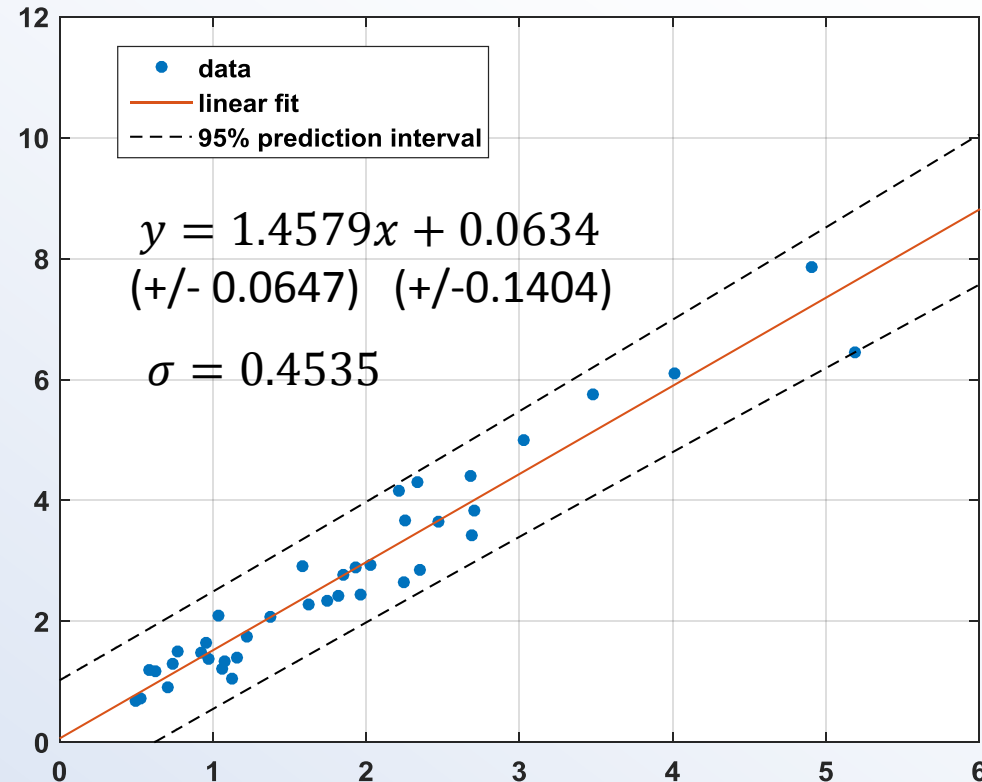


- Estimate from **blue series** what would **red series** be if it had been measured in the past.

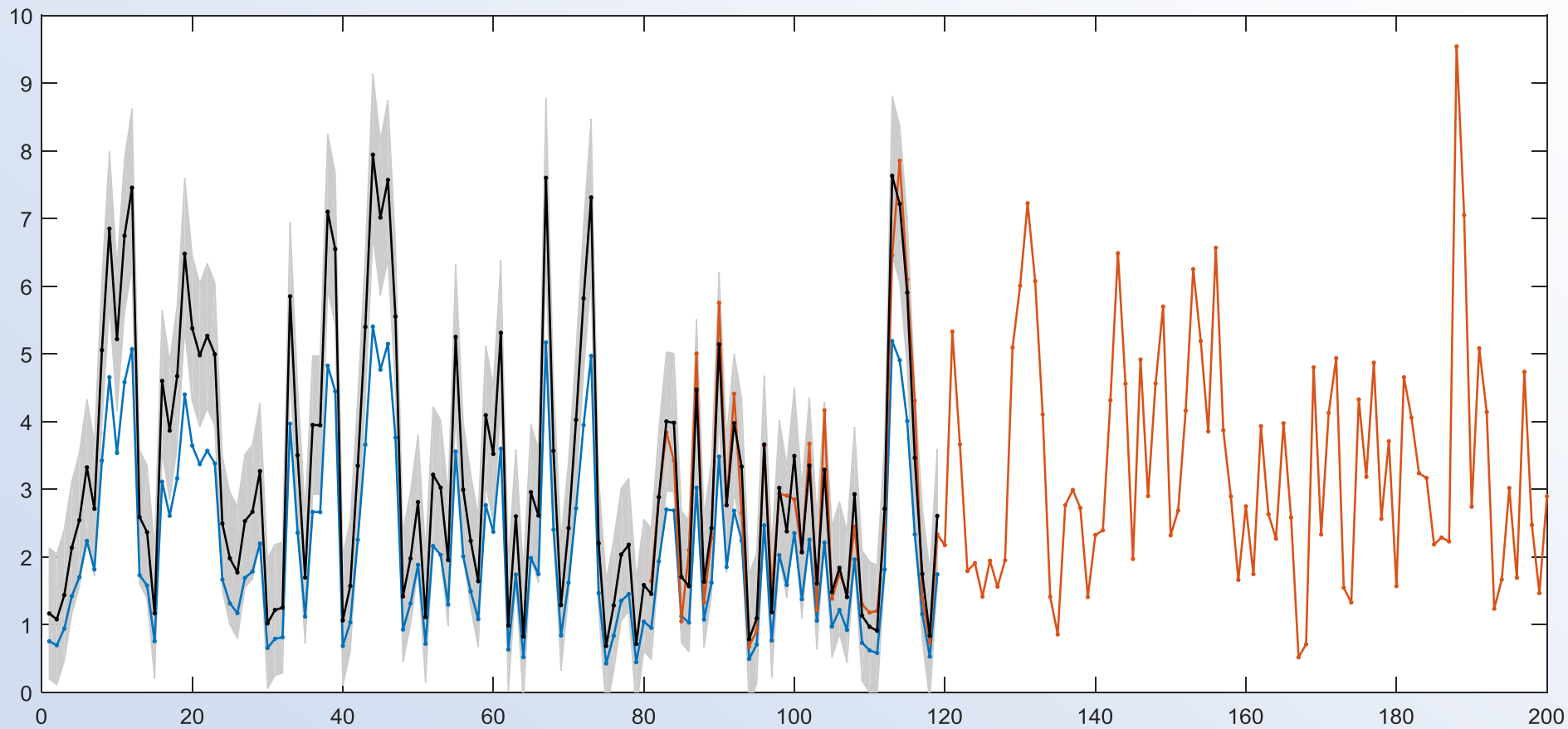


- However, important to estimate error!
- Regression
$$y = \beta_0 + \beta_1 x + \varepsilon$$
- $\varepsilon = N(0, \sigma)$
- Since σ is not known sampling distribution of ε is Student T-distr. With $n-2$ DF
- 95% prediction interval:

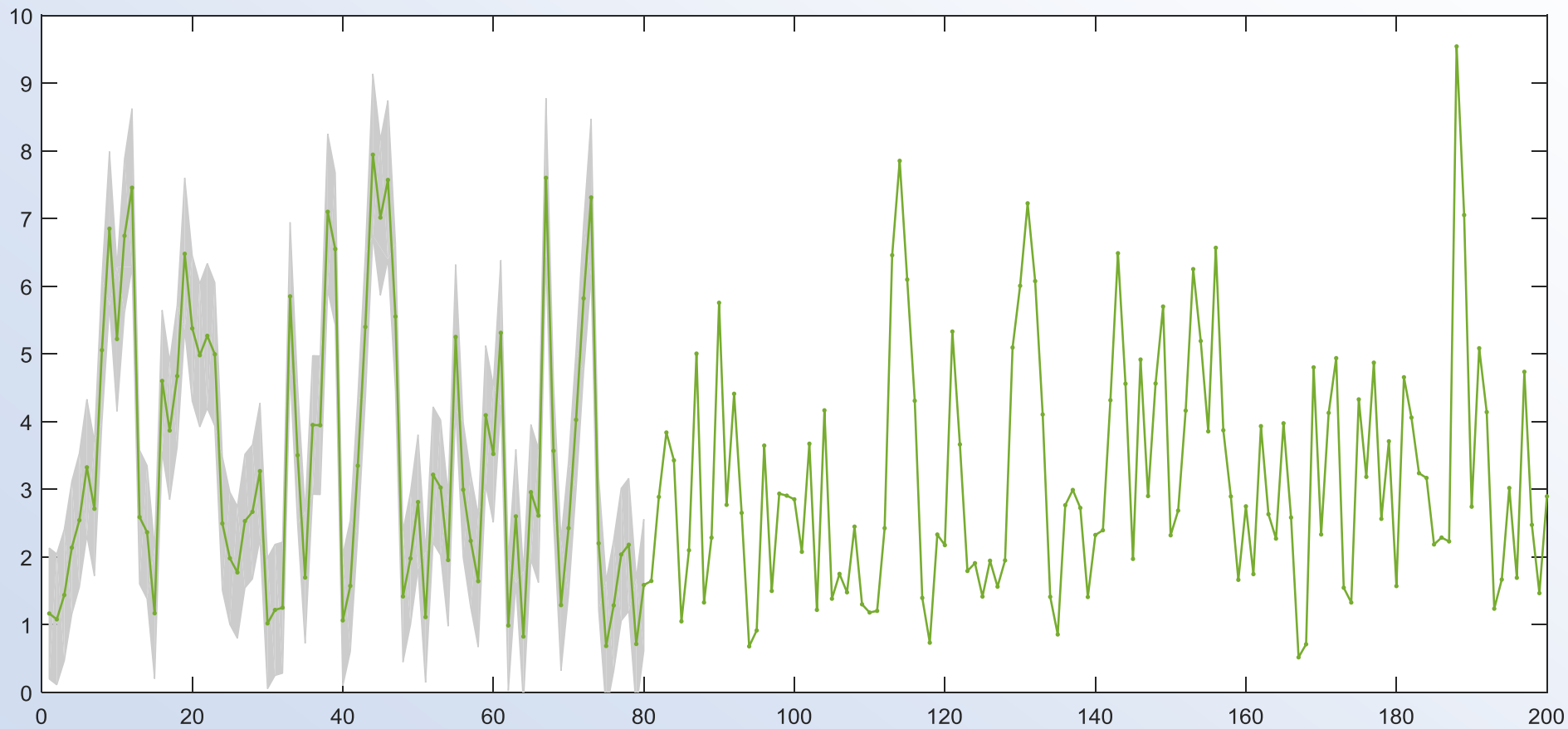
$$\Delta y = t_{0.975, n-2} \sqrt{\text{var}(\beta_0) + x^2 \text{var}(\beta_1) + \sigma^2}$$



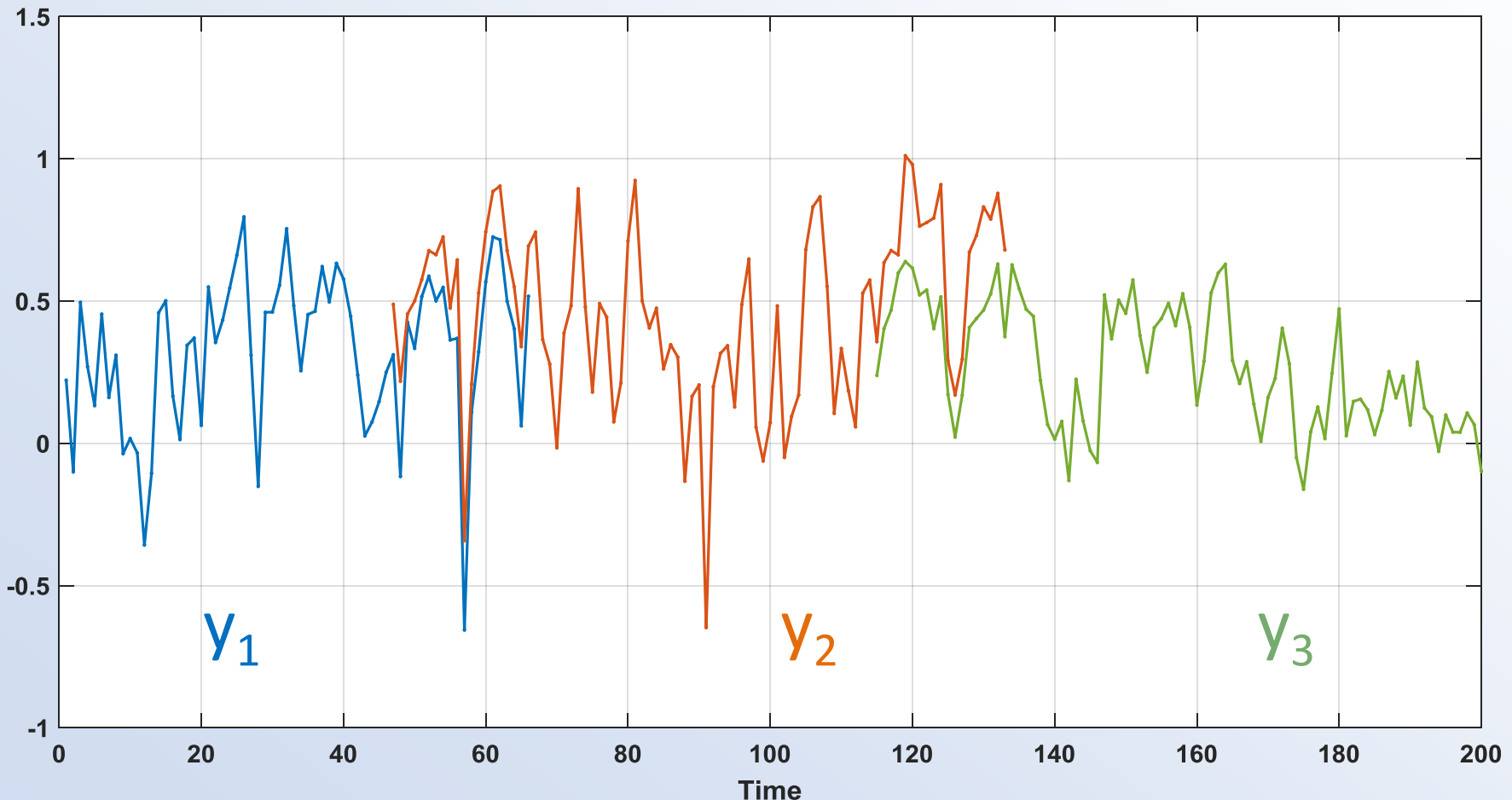
- Composited data is now black and red series



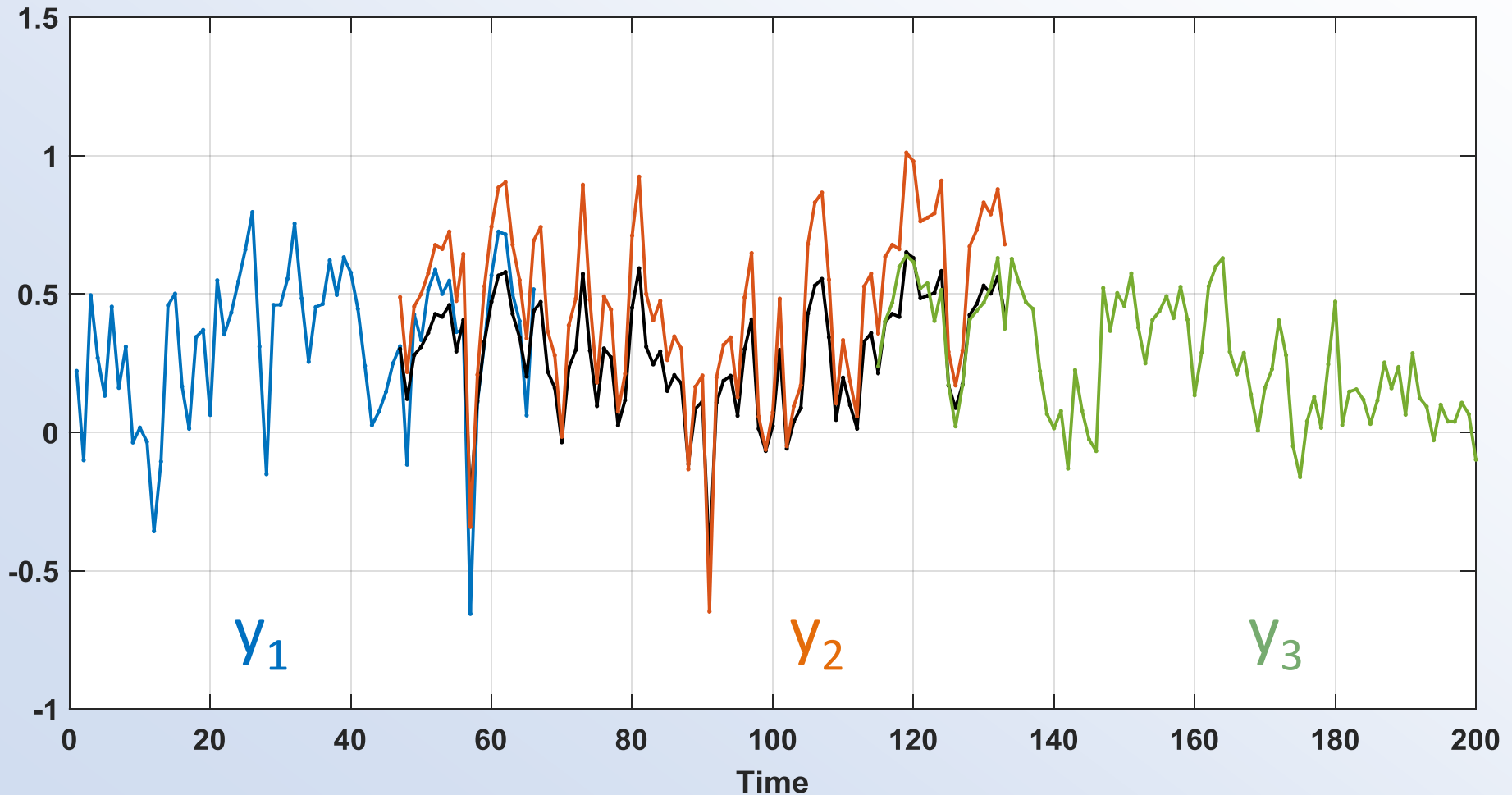
- Composited data is now **green** series



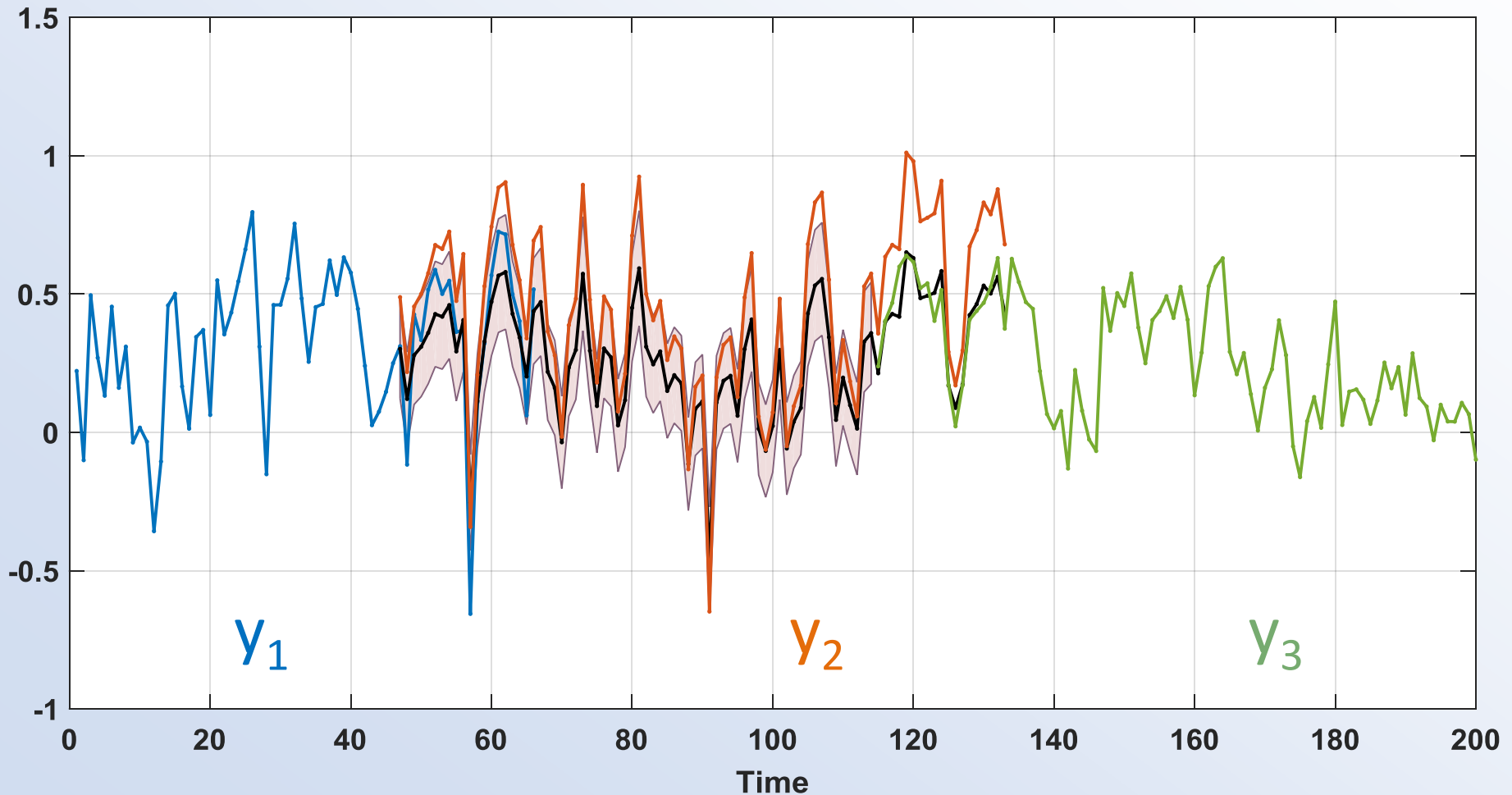
- The above method can be generalized to multiple overlapping datasets. **We select y_3 as the basis**



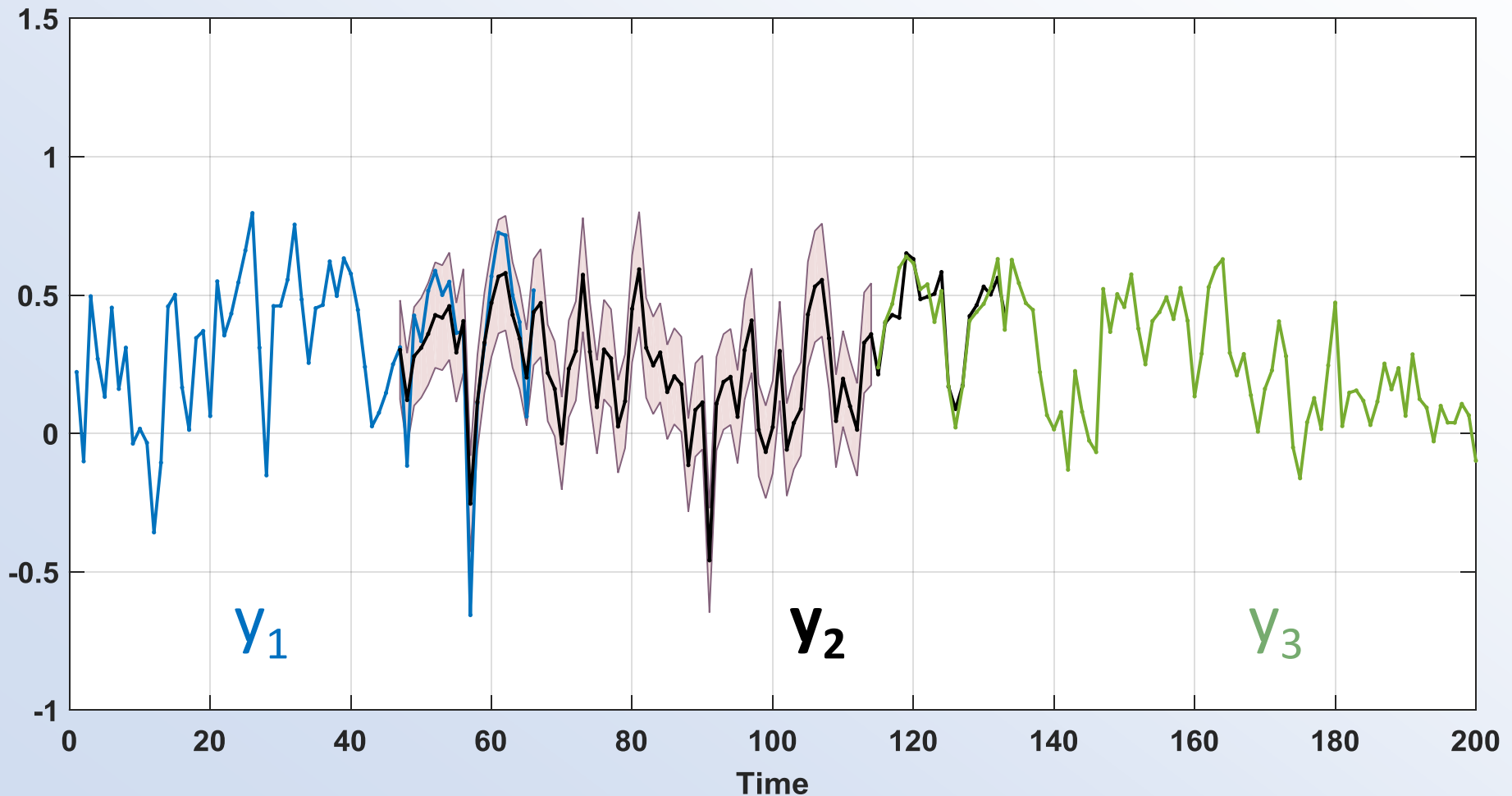
- First normalize y_2 to level of y_3



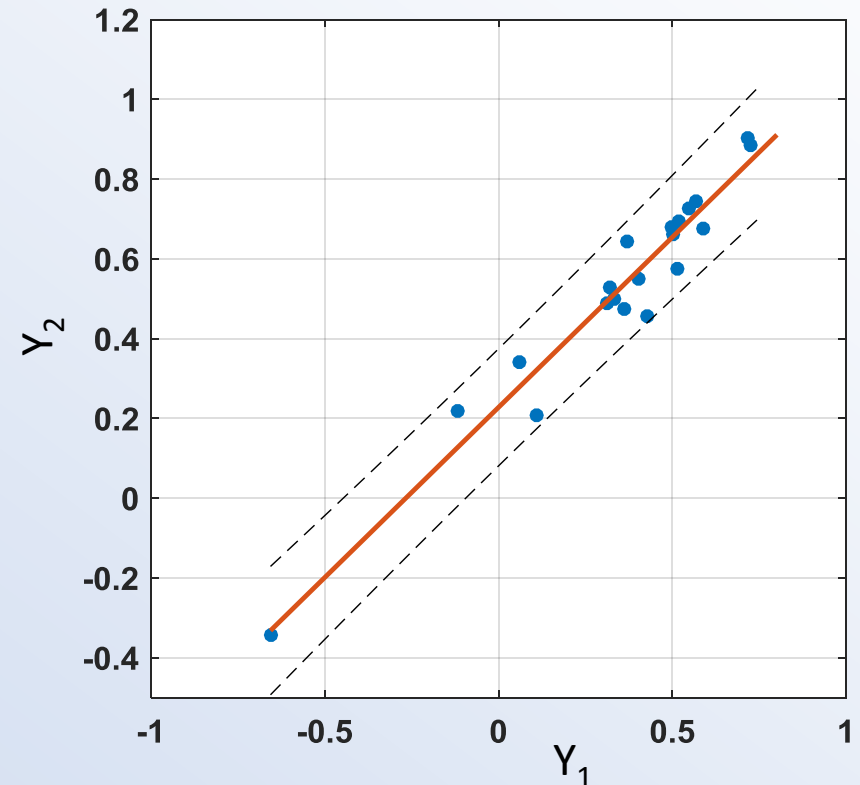
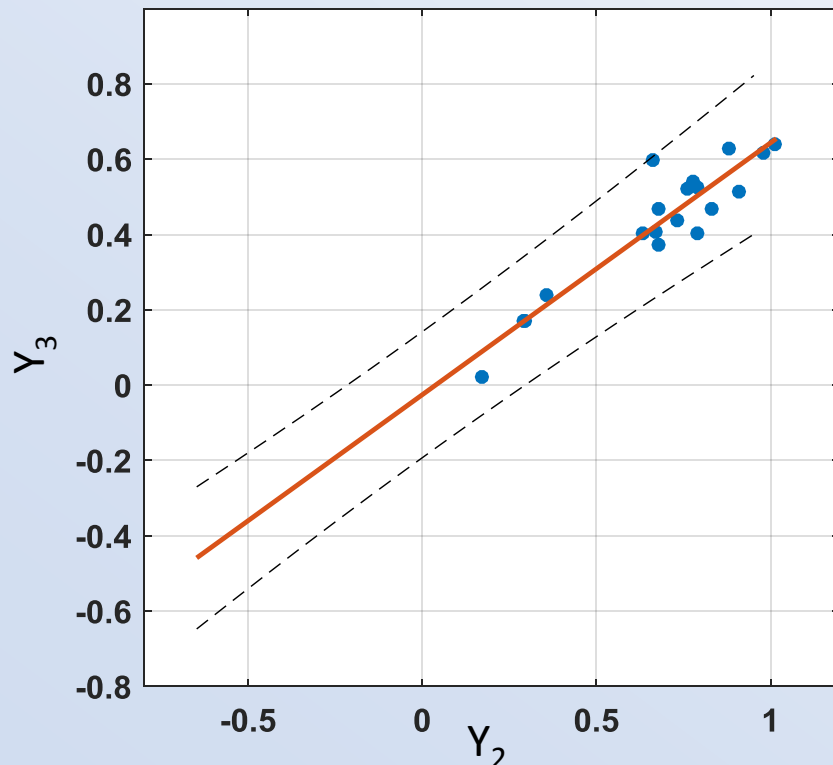
- First normalize y_2 to level of y_3
- This normalization has some **statistical error**



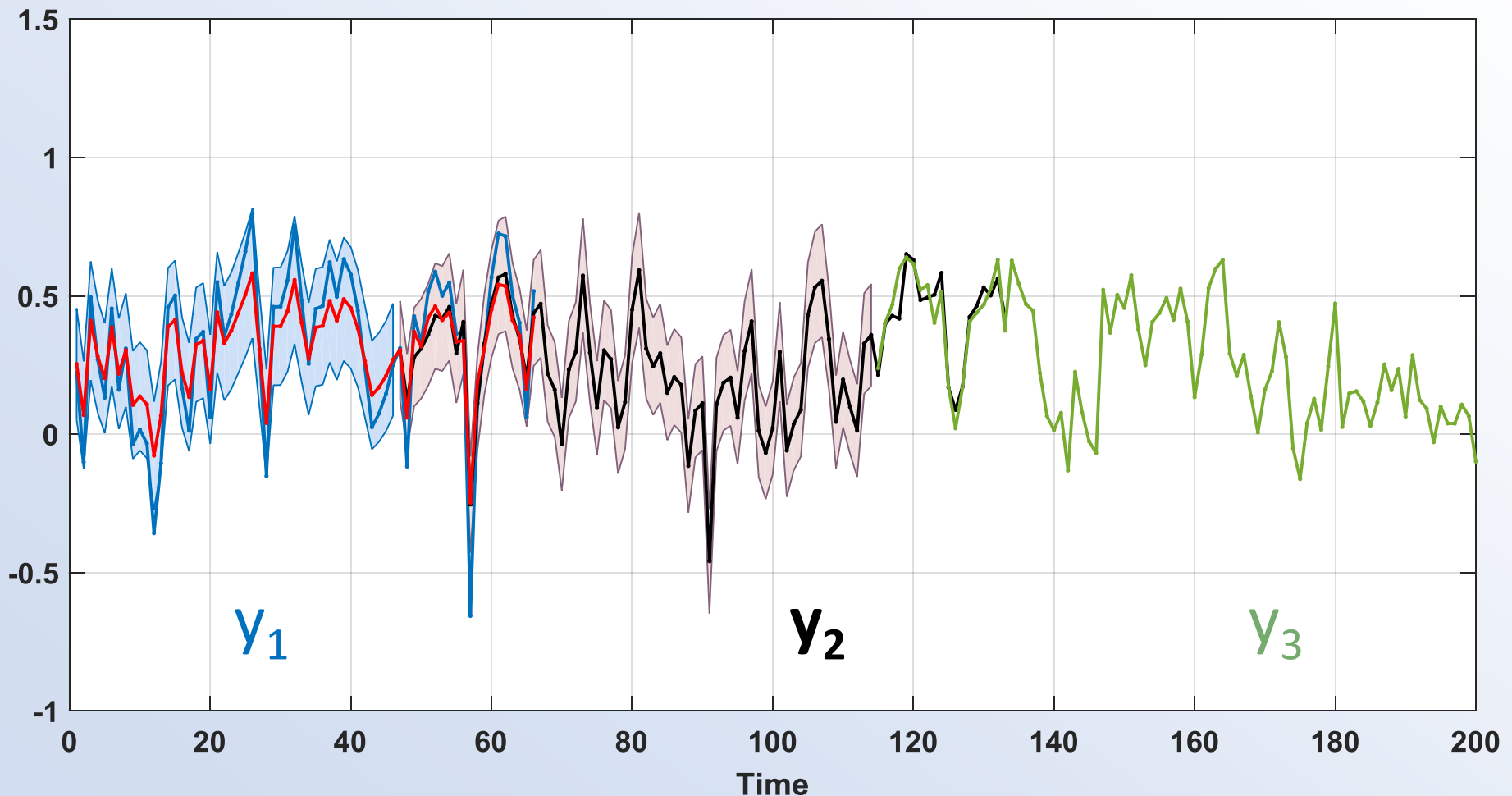
- Next normalize y_1 to new y_2



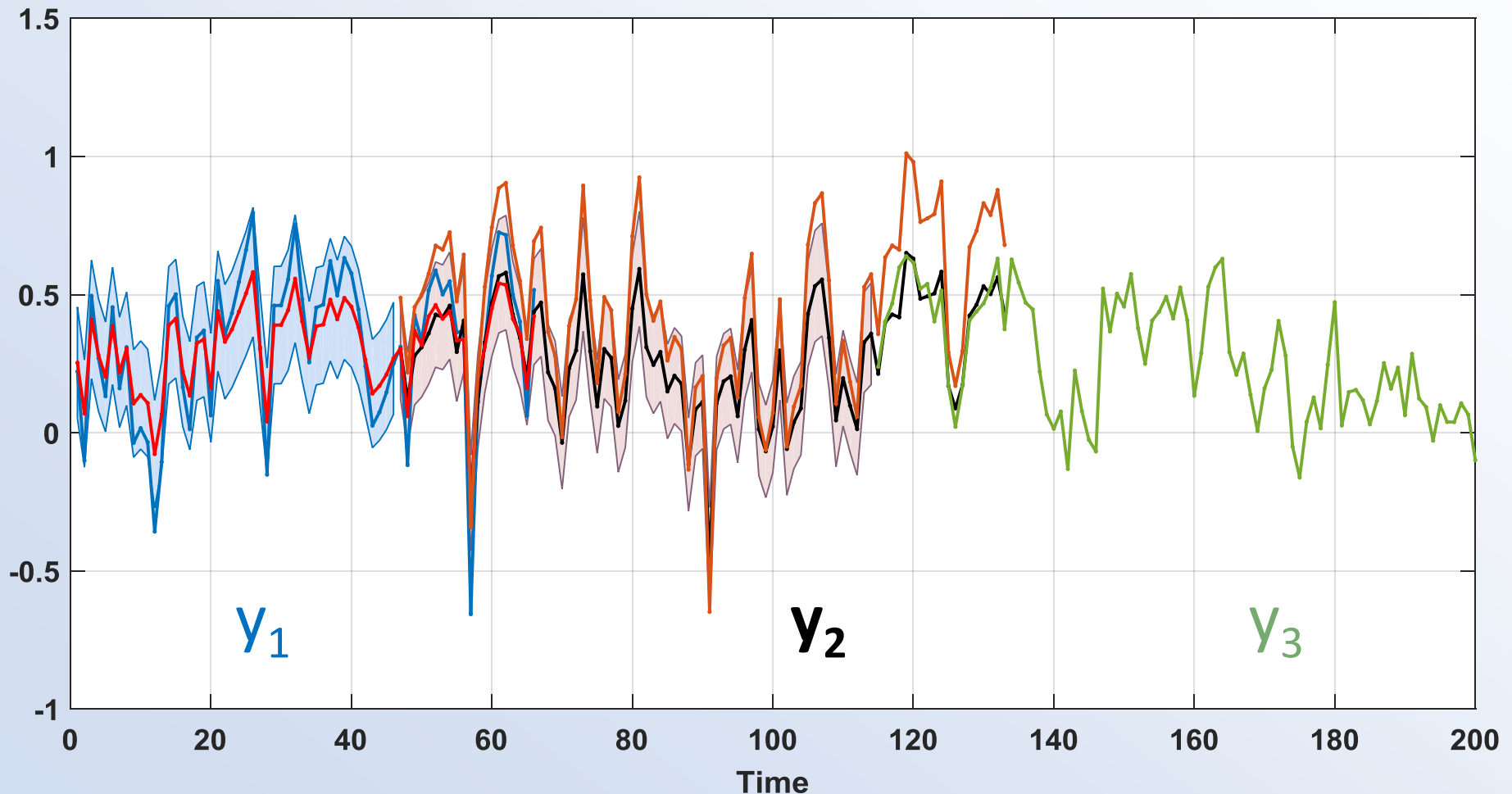
- From the two regressions we have
- 1st norm: $y_3 = \beta_1 + \beta_2 y_2$
- 2nd norm: $y_2 = \alpha_1 + \alpha_2 y_1$
- 2nd norm: $y_3 = \beta_1 + \beta_2(\alpha_1 + \alpha_2 y_1)$
- $var(y_2) = var(\alpha_1) + y_1^2 var(\alpha_2) + var(\varepsilon_1)$
- $var(y_3) = var(\beta_1) + y_2^2 var(\beta_2) + var(y_2)\beta_1^2 + var(\beta_2)var(y_2) + var(\varepsilon_2)$
- 95% Prediction interval is given by:
- $\Delta y = t_{0.975, n-4} \sqrt{var(y_3)}$



- Next normalize y_1 to new y_2
- And estimate error limits

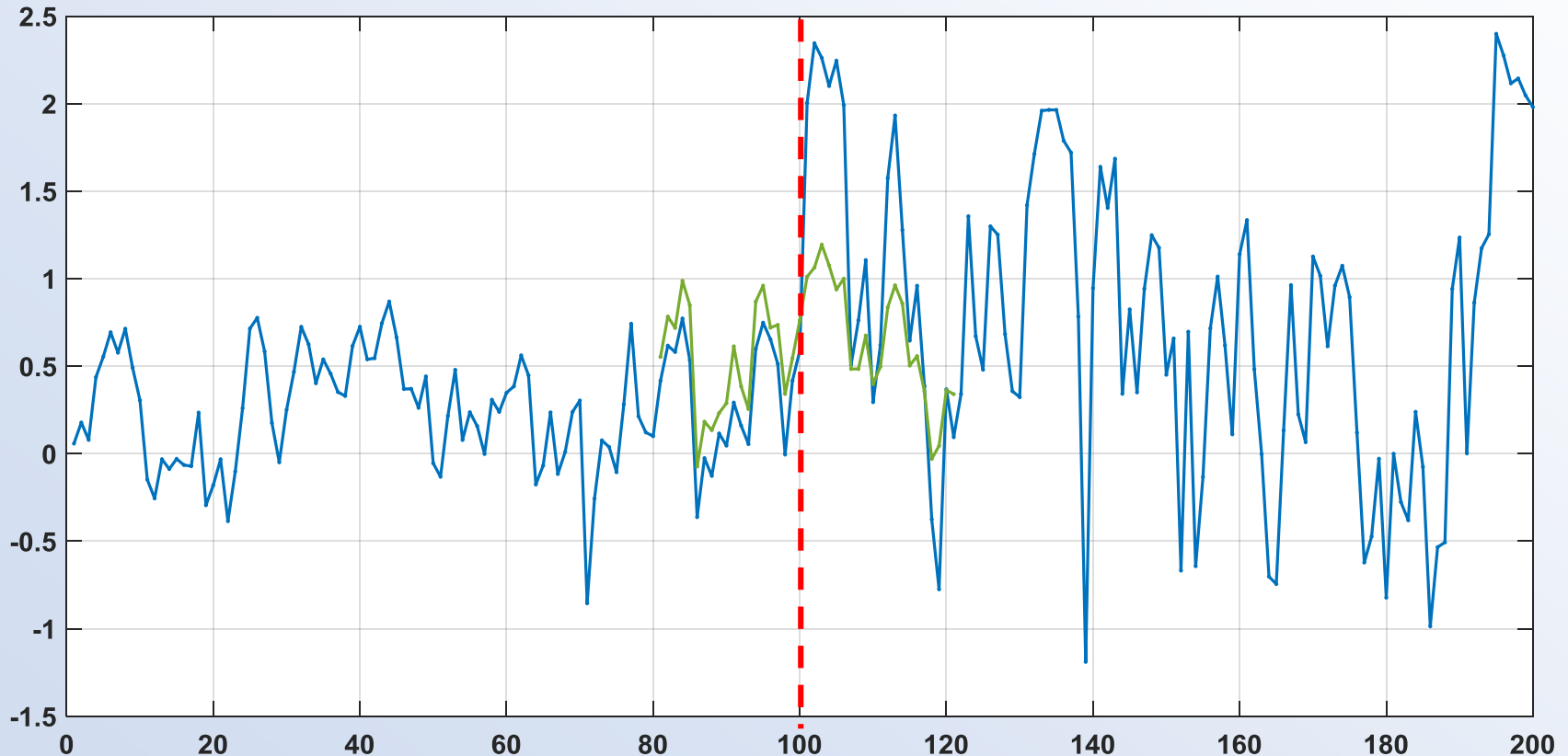


- Finally we have the composited data series and its error limits.
- Note: **Errors are accumulated**. In this example the errors in the first time series are larger than the difference between the original and the composite!



Removing inhomogeneity with a reference series

- Often a single long data series can be inhomogeneous due to changes in instrument settings, location etc.
- Sudden changes can be calibrated with a reference time series, which is homogeneous over the same time interval



- x_1 =data before jump, x_2 =data after the jump, y = reference series

- After the jump:

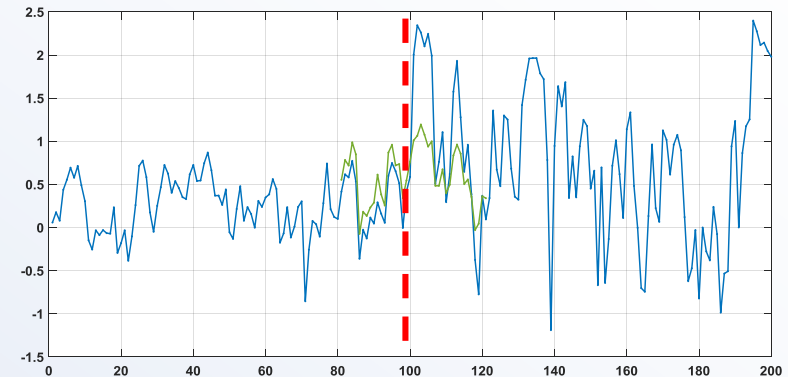
- $x_2 = \beta_1 + \beta_2 y + \varepsilon_2$ n_2 points

- Before the jump:

- $y = \alpha_1 + \alpha_2 x_1 + \varepsilon_1$ n_1 points

- Estimate what x_2 would be BEFORE the jump

- $\rightarrow x_2 = \beta_1 + \beta_2(\alpha_1 + \alpha_2 x_1 + \varepsilon_1) + \varepsilon_2$



- From summation of variance formula we get for the time period before the jump

- $var(y) = var(\alpha_1) + x_1^2 var(\alpha_2) + var(\varepsilon_1)$

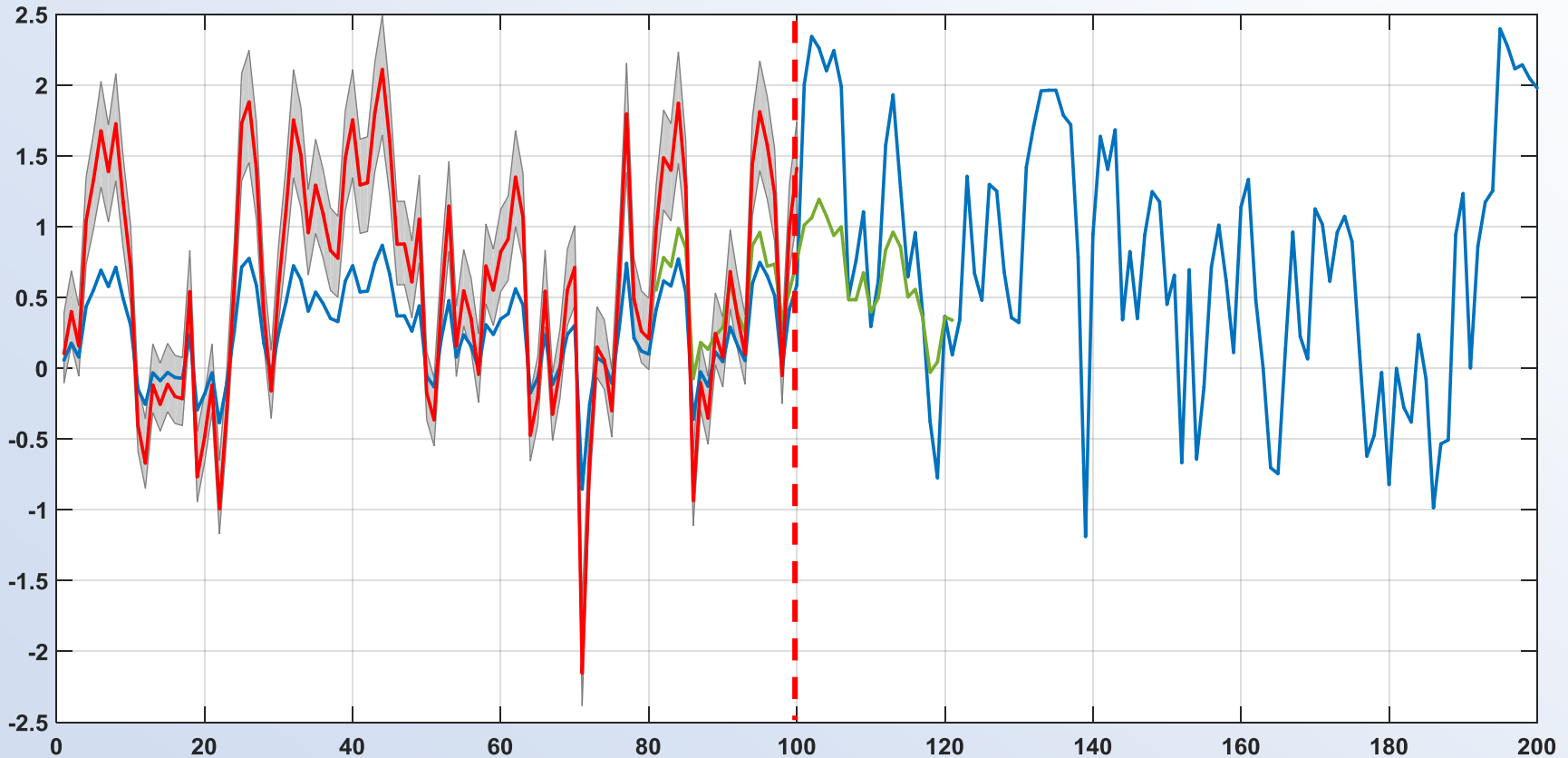
- $var(x_2) = var(\beta_1) + y^2 var(\beta_2) + \beta_2^2 var(y) + var(\beta_2)var(y) + var(\varepsilon_2)$

- And the 95% prediction interval would be

- $\Delta x_2 = t_{0.975, n_1+n_2-4} \sqrt{var(x_2)}$

Removing inhomogeneity with a reference series

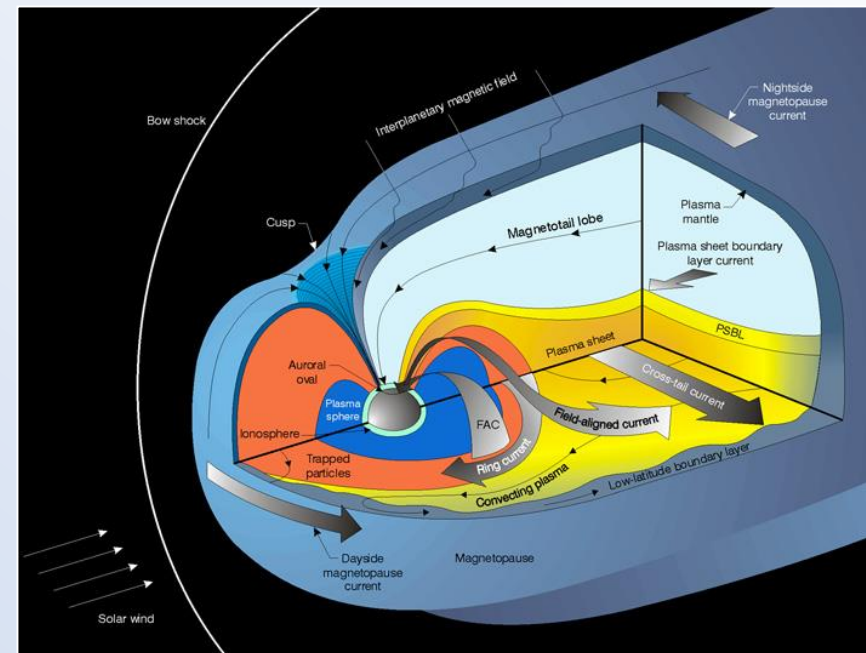
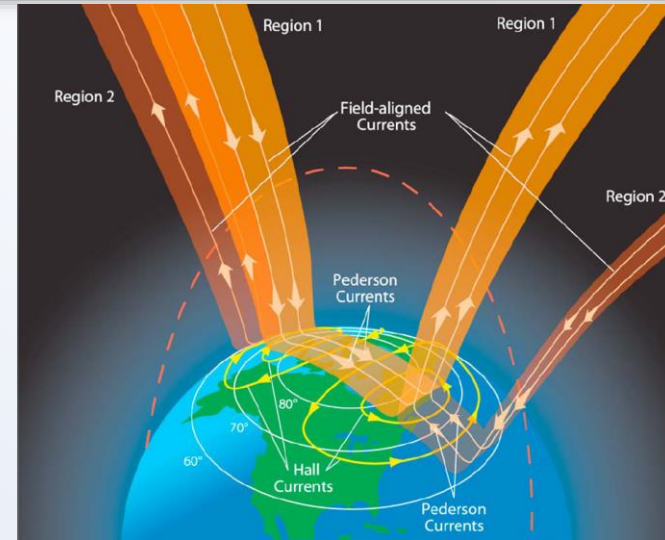
- Here the inhomogeneity has been corrected



Example 1.

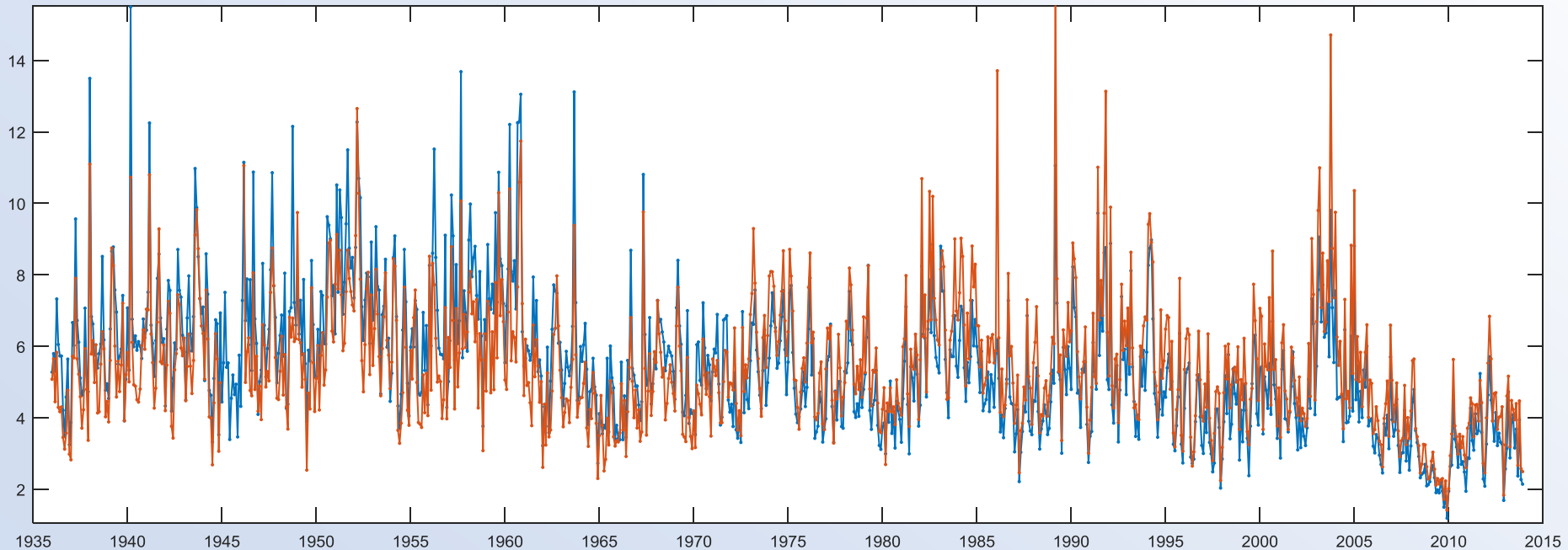
Geomagnetic indices

- Geomagnetic activity
 \Leftrightarrow Variations of magnetic field on ground
- Results from variations of electric currents in magnetosphere and ionosphere

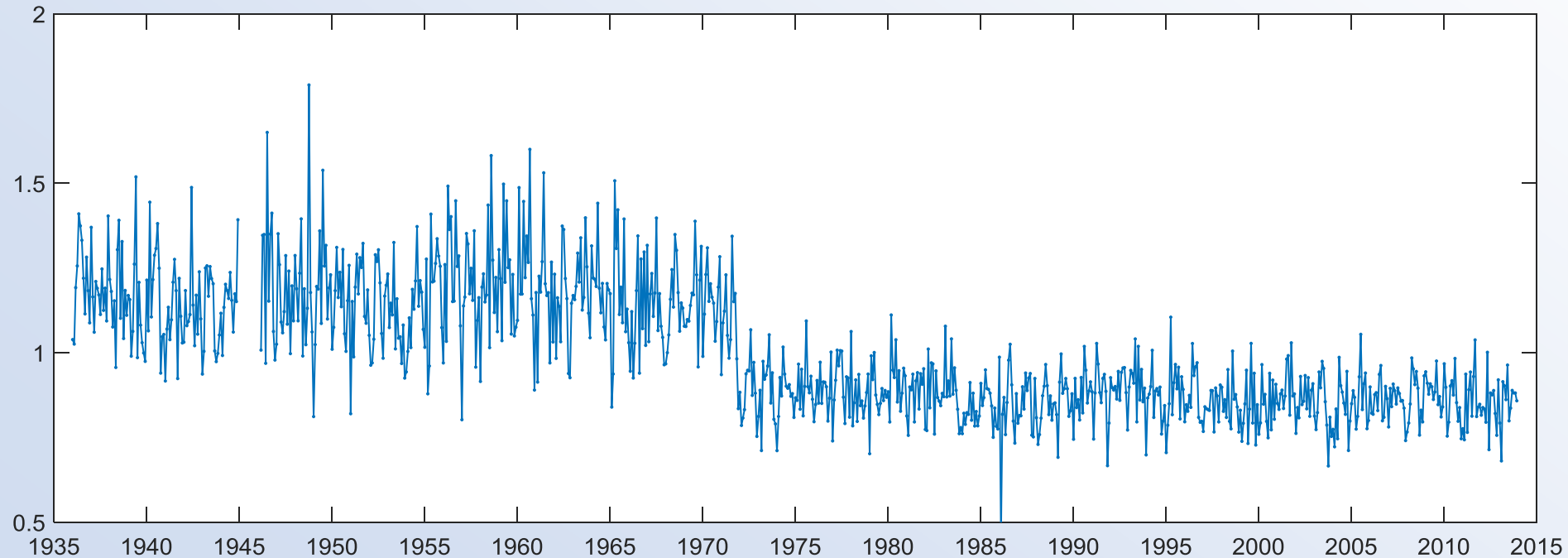


- IHV indices from two different mid-latitude stations (NGK/Germany, CLF/France)
- CLF changed from spot sampling to hourly means in 1972

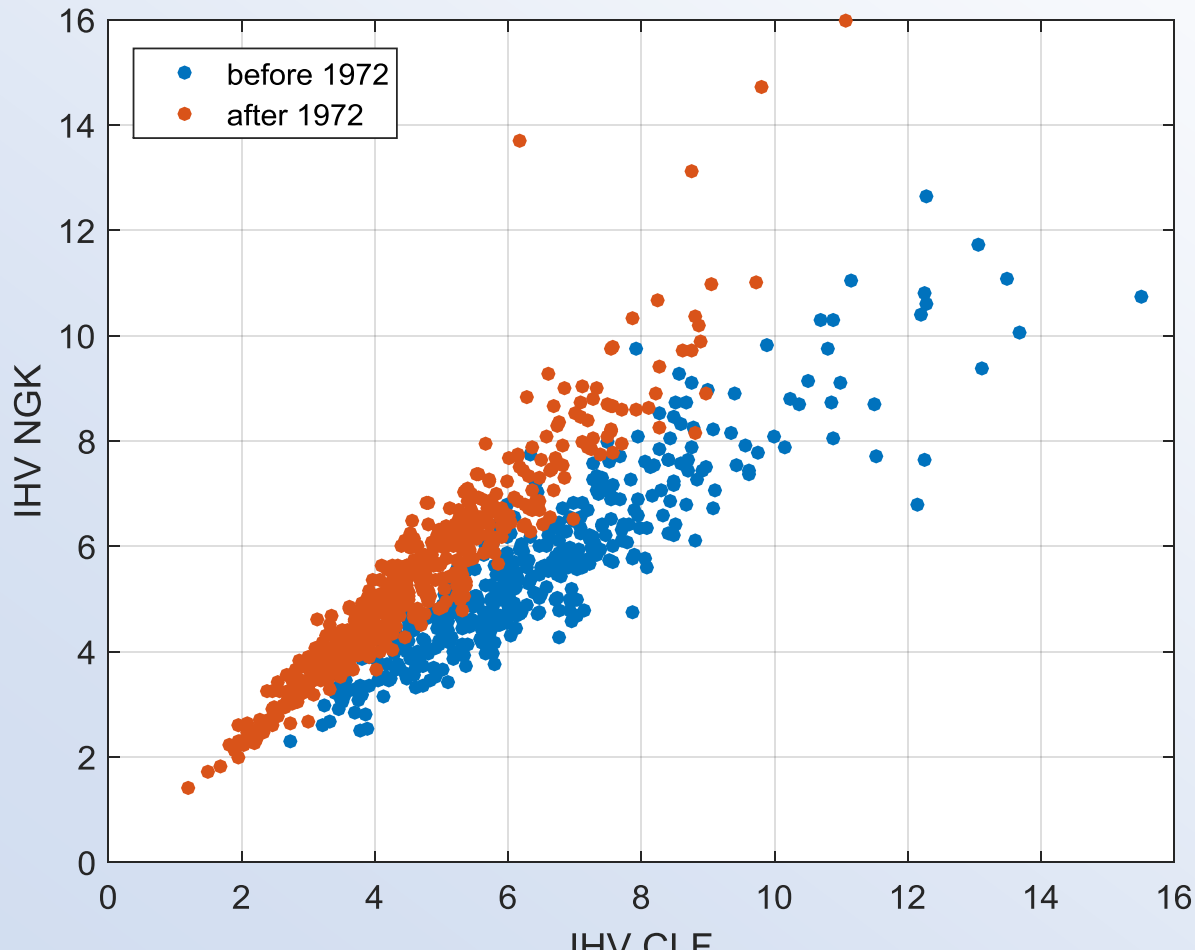
$$IHV = \frac{1}{6} \sum_{t=21LT}^{02LT} |H(t+1) - H(t)|$$



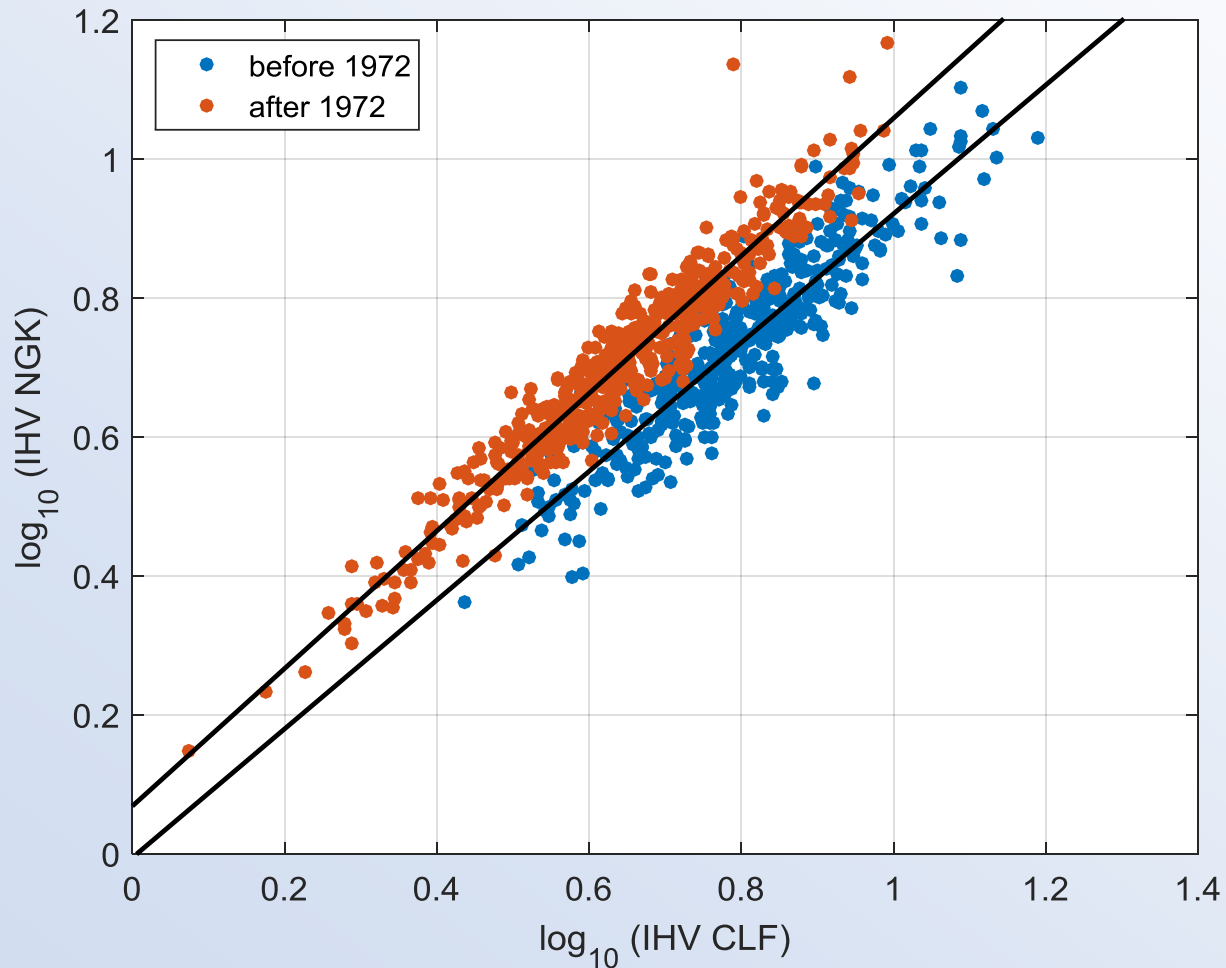
- Effect of sampling change is clearly seen in CLF/NGK ratio
- Before 1972 CLF sees systematically larger values than NGK compared to period after 1972
- Note! Ratio after 1972 is not 1 → Real difference in the indices.
- We are not trying to make the indices identical!!



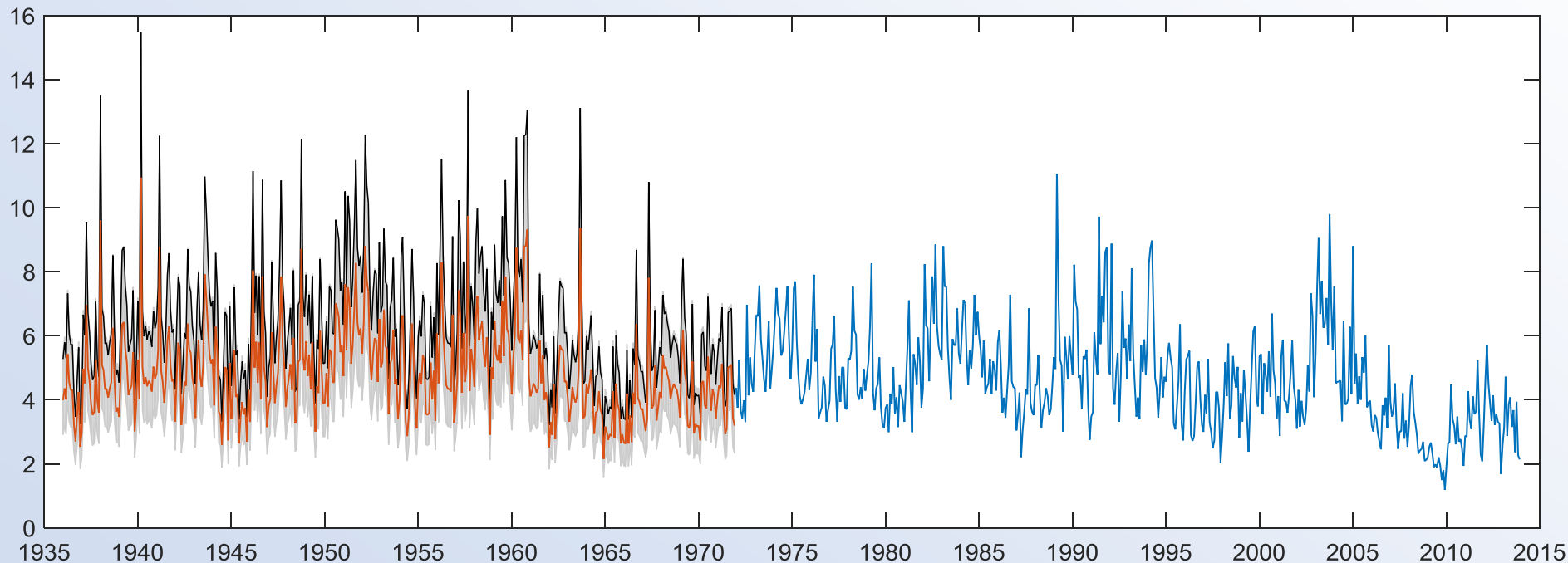
- Clear difference in two time periods
- But linear fit is not good (not a constant variance)



- Log-log scale much better



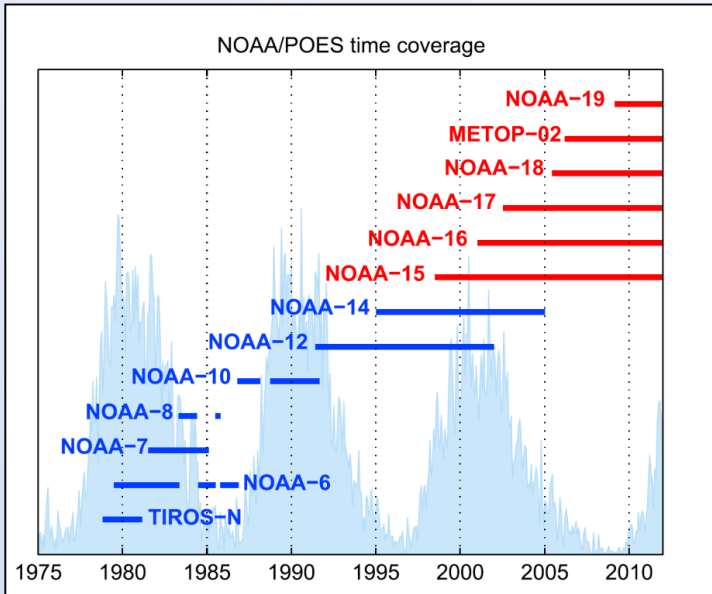
- After calibration the data before 1972 is in correct level and we have also uncertainty limits



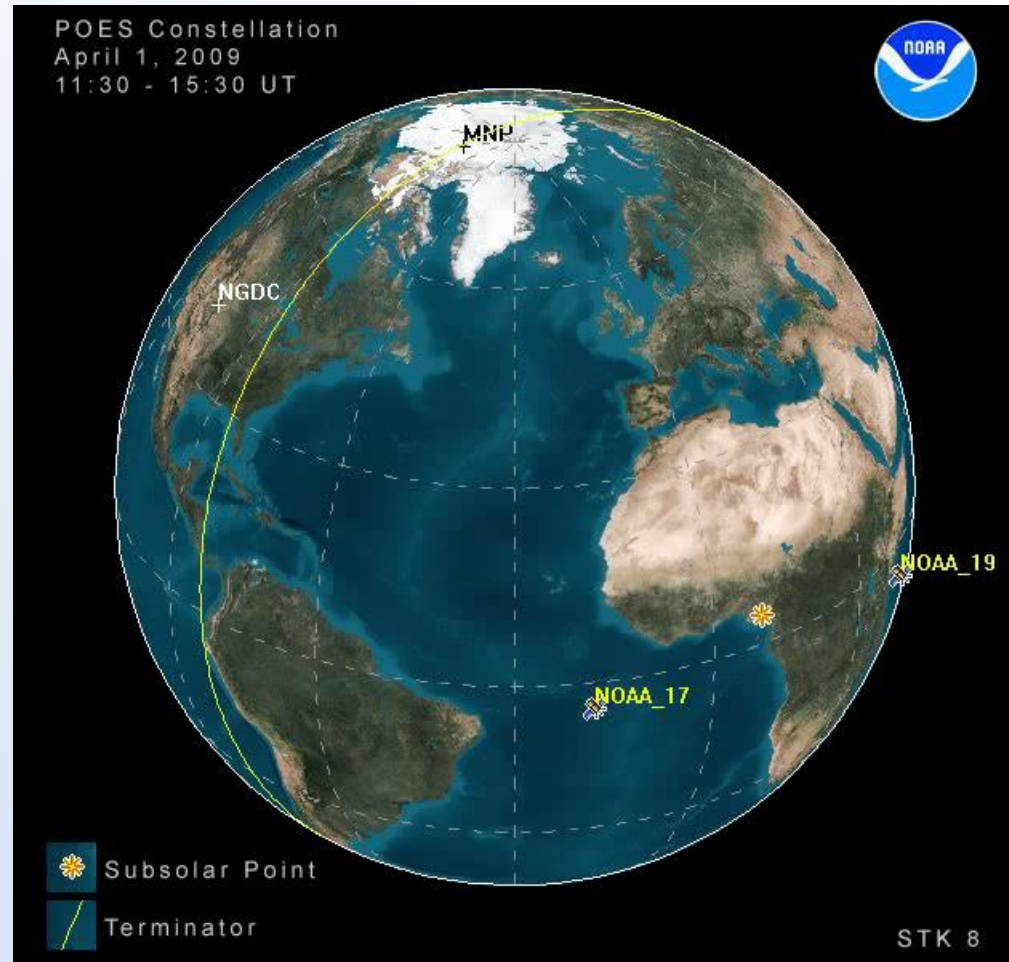
- In many cases **inhomogeneity** of a measurement series **changes with time**
- Examples:
 - Instrument efficiency changes continuously with time
 - Instrument degrades with time
- Often these cannot be corrected by comparing with some other time series at one interval of time
- **→ Need to understand the cause of changing data homogeneity and how to compensate it**

Example 2.

Particle data from satellites

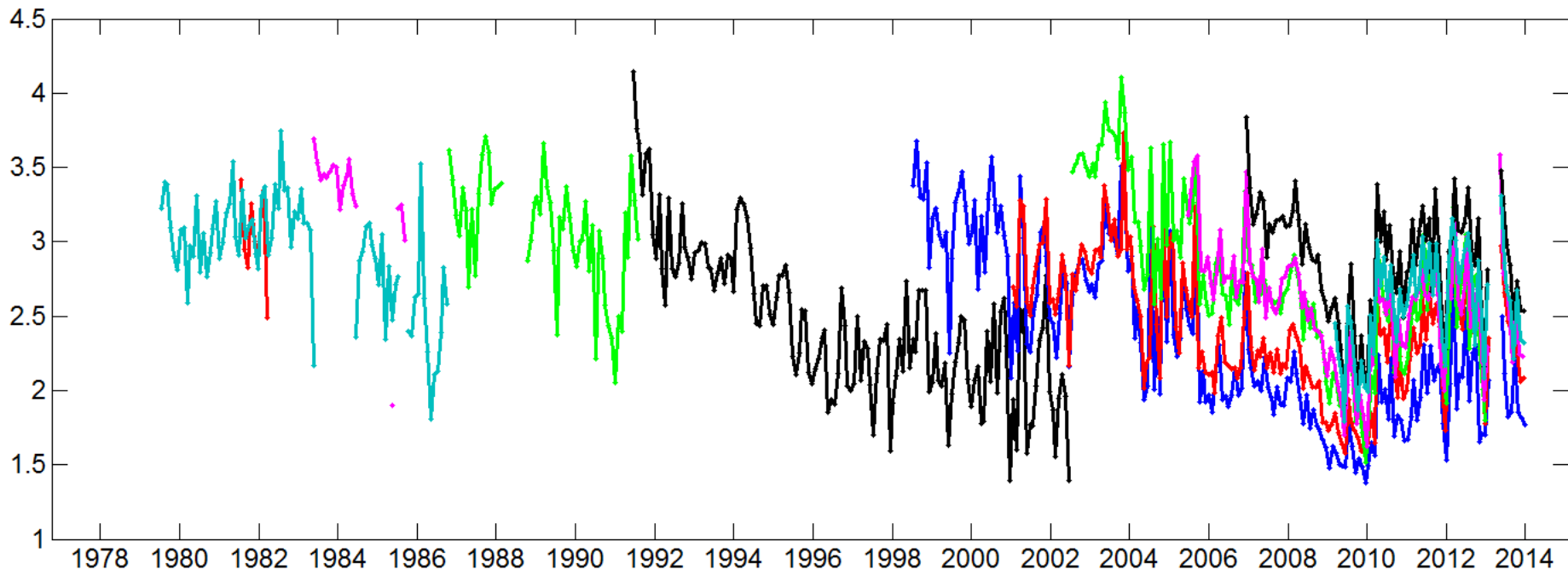


- The dataset has been plagued by significant problems related to the MEPED instrument.

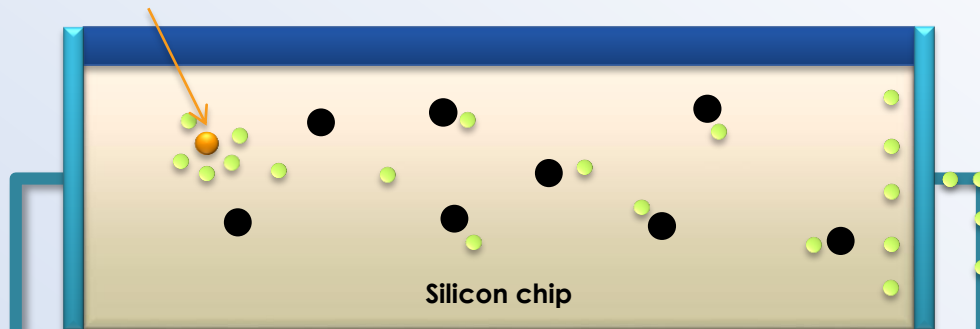


Raw proton data is obviously erroneous

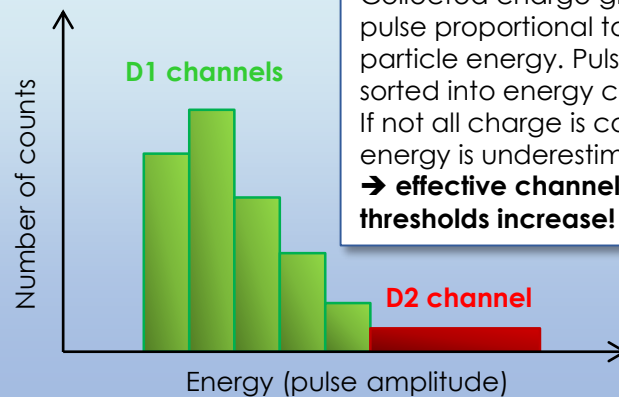
- Steady decrease in all satellites → degradation of instrument
- Large spread between simultaneous satellites
- Simple stitching of overlapping data series does not work.



Zoom-up of a single silicon detector chip

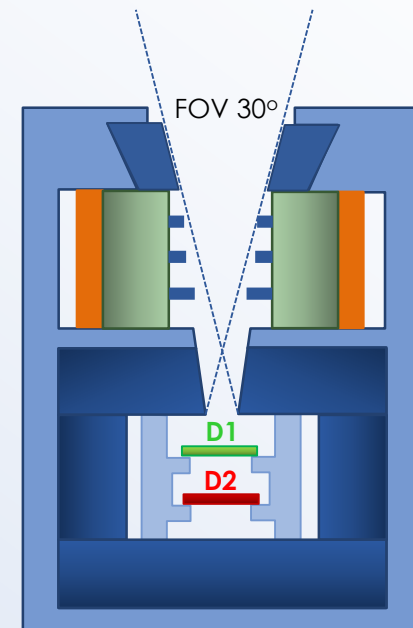


Charge collecting electronics and pulse height analyzer



Collected charge gives a pulse proportional to particle energy. Pulses are sorted into energy channels. If not all charge is collected, energy is underestimated
→ effective channel thresholds increase!

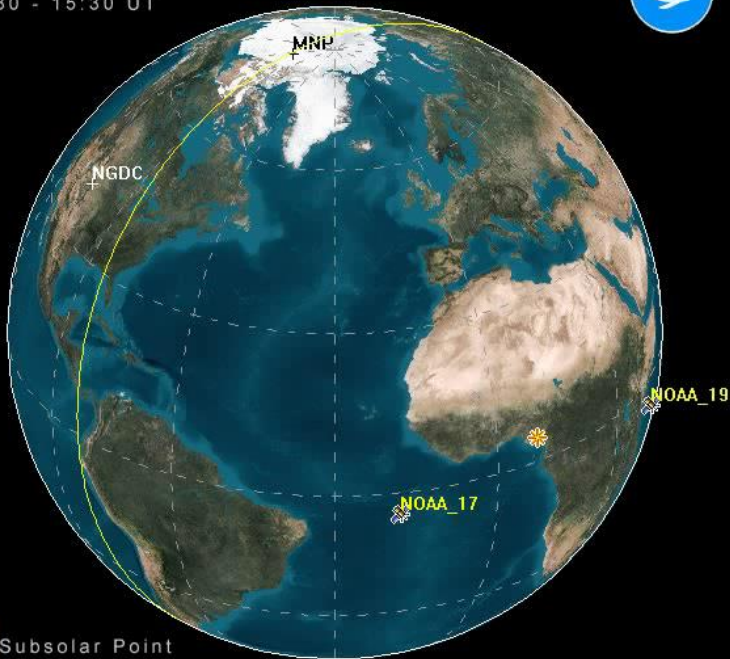
Cross-cut of the MEPED proton instrument





Anti-coincidence logic between **front detector (D1)** and **back detector (D2)**
→ a noise(false) count in D2 erases a coincident real count from D1.

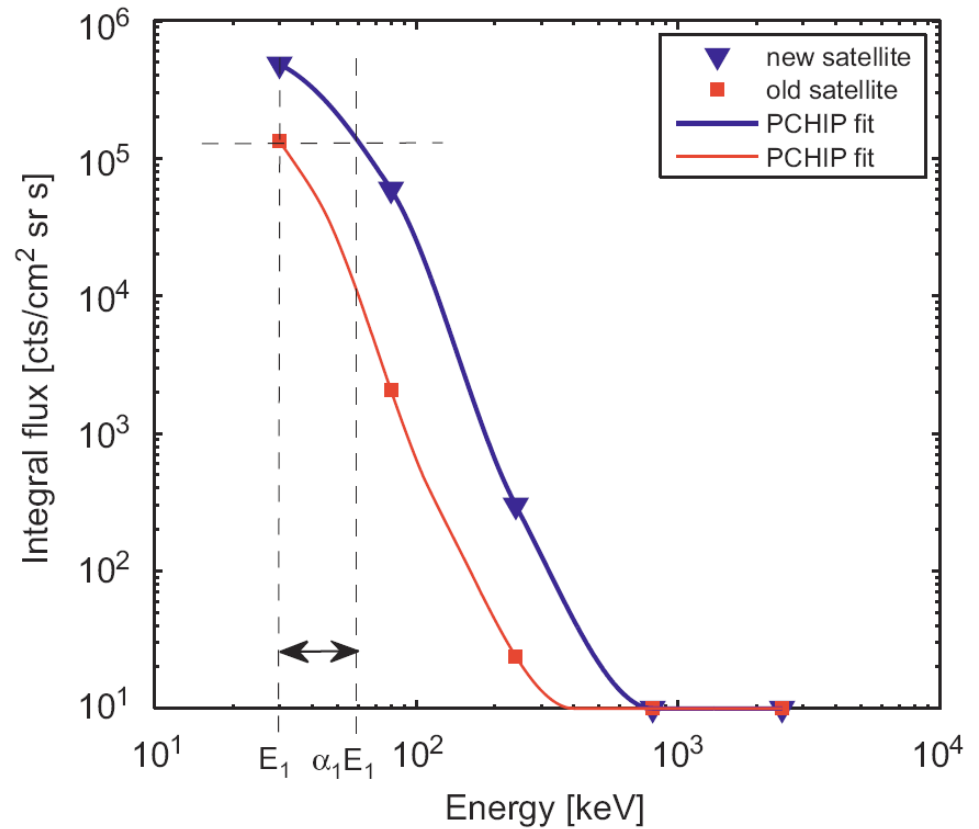
Estimate instrument energy thresholds by comparing an old satellite to new

POES Constellation
April 1, 2009
11:30 - 15:30 UT

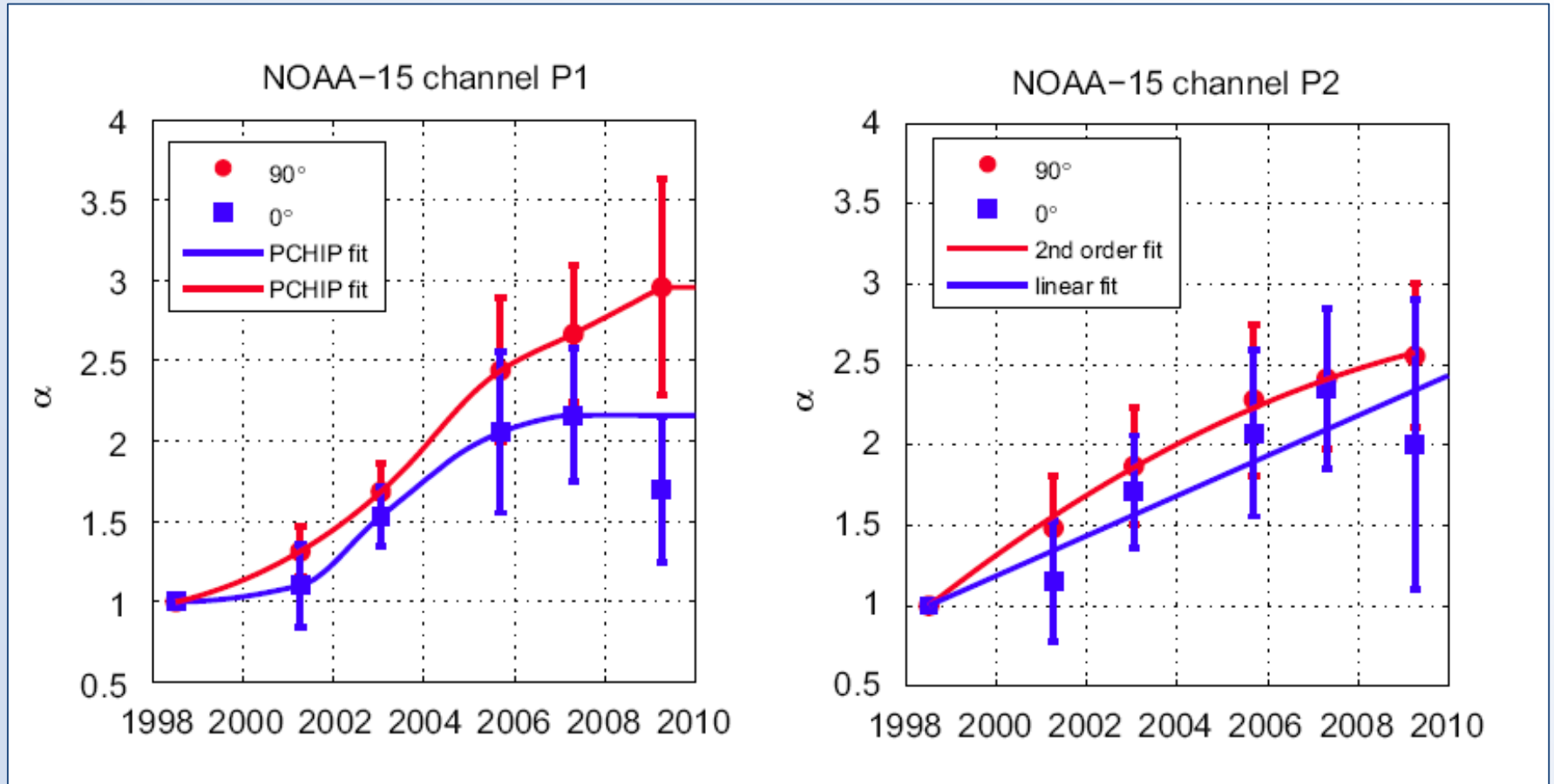


 Subsolar Point
 Terminator

STK 8

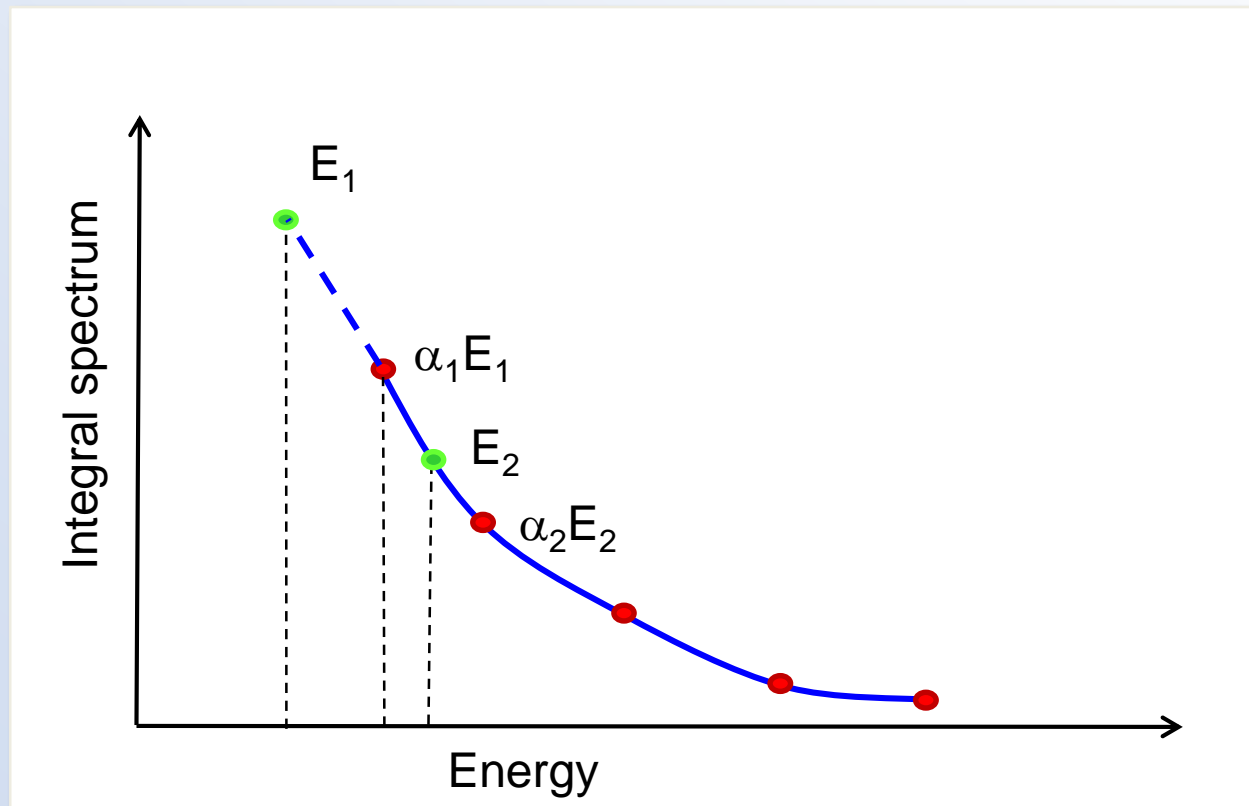


We get real energy thresholds as a function of time



Compute fluxes at nominal energies from spectrum fitted to actual energies

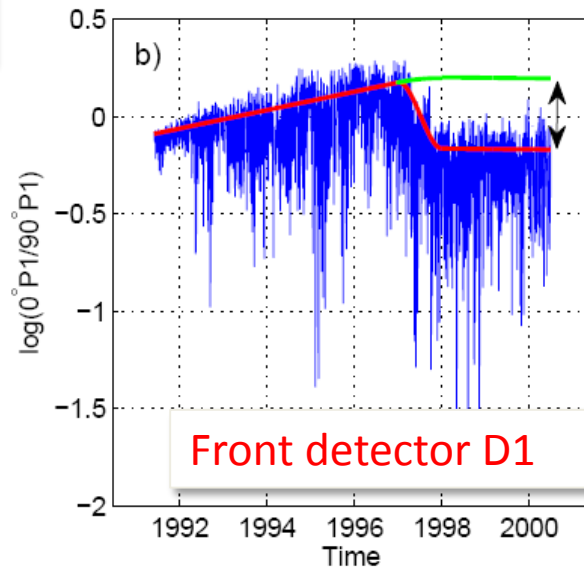
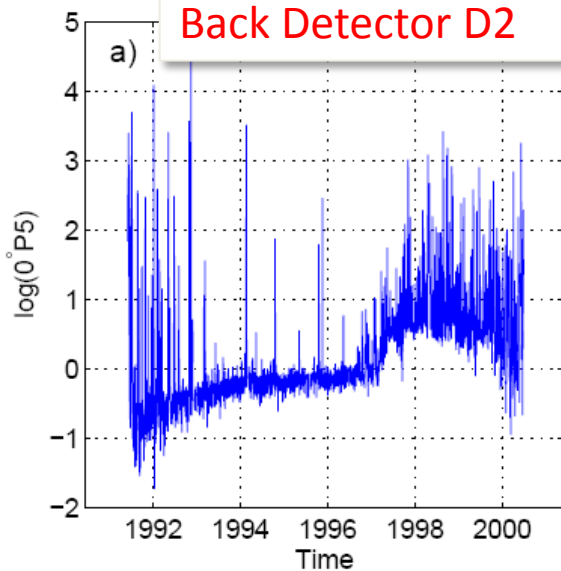
- Note a caveat! The P1 energy channel cannot be reliably corrected, since it requires extrapolation to energies lower than the instrument can measure at the time.



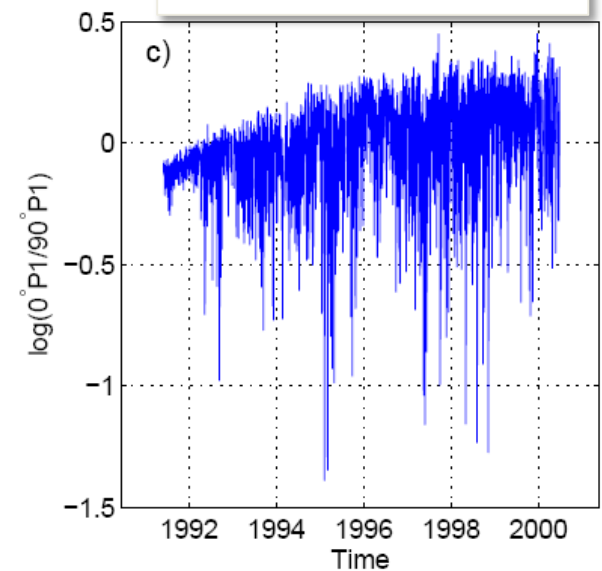
Electronic back detector noise in NOAA-08 and NOAA-12

- Back and front detectors work in anti-coincidence logic
- **False counts (noise) in back detector erase real counts from front detector**
- Modify the front detector baseline to correct for the noise

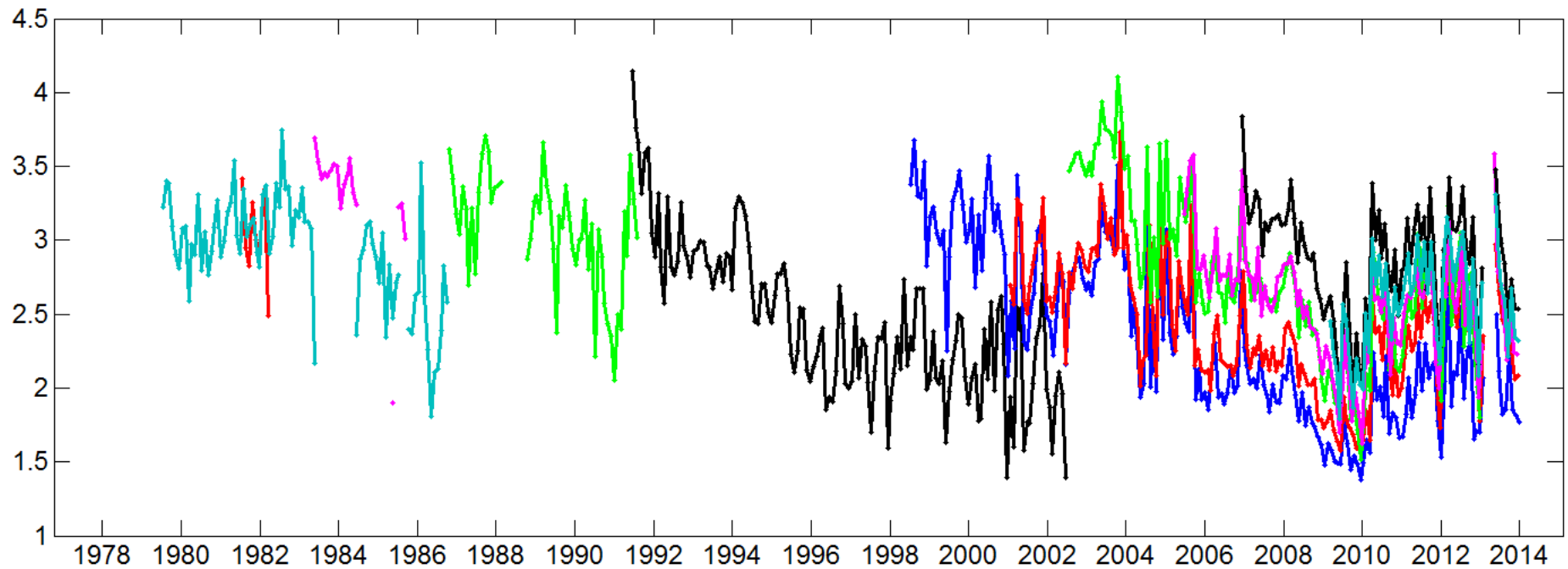
a) Back Detector D2



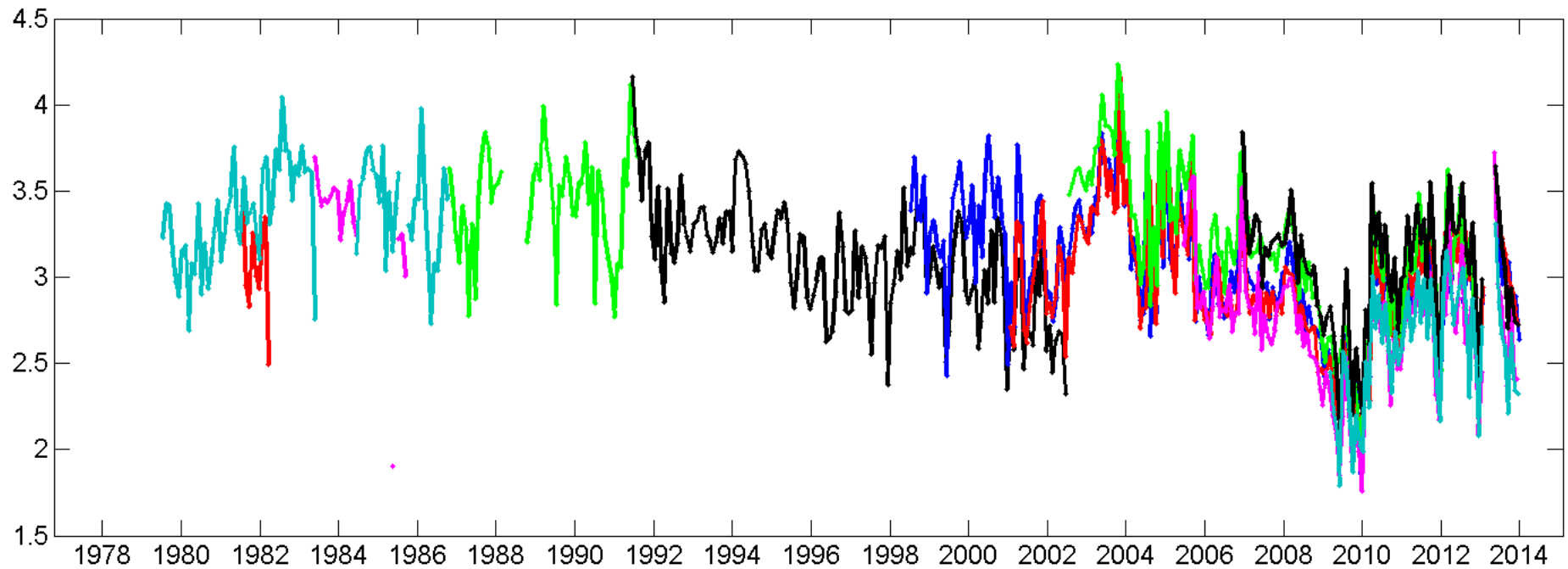
Front detector D1



- Steady decrease in all satellites → degradation of instrument
- Large spread between simultaneous satellites

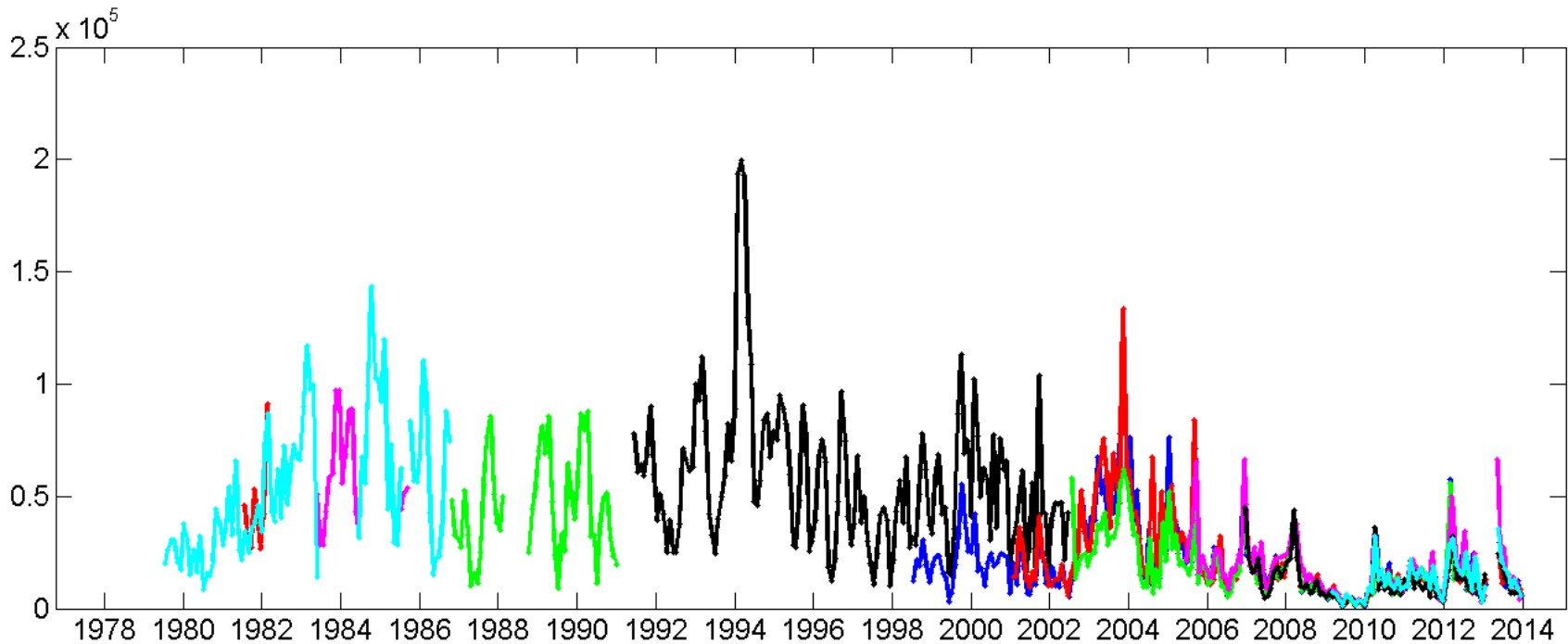


- Continuous series from different satellites
- Satellite differences greatly reduced



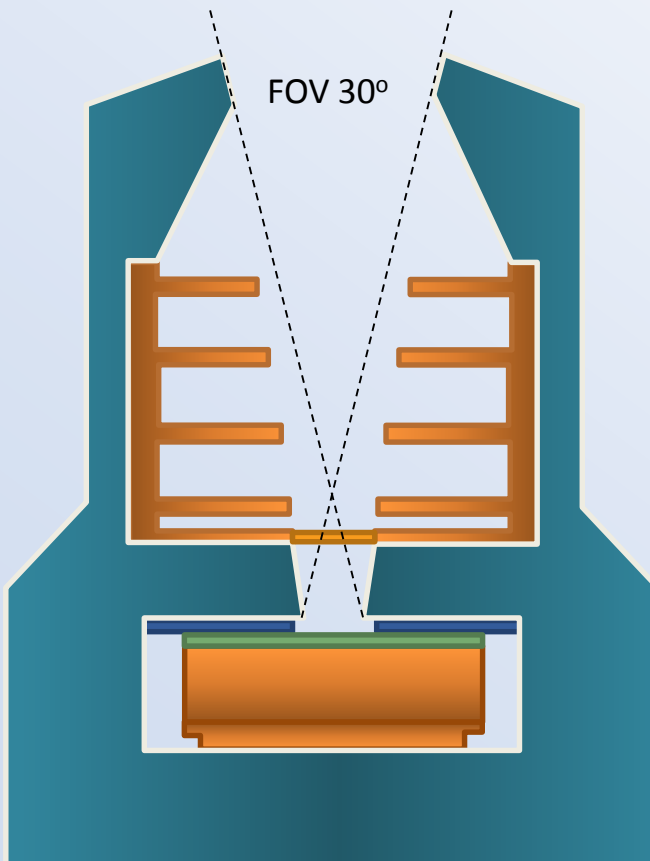
Electrons > 30 keV uncorrected

- Newer SEM-2 satellites (since mid-1998) systematically smaller than SEM-1 satellites



- Satellites up to NOAA-14 carry an **older SEM-1 instrument**. Starting with NOAA-15 (mid 1998) the satellites carry **a newer SEM-2 instrument**.
- Problems in MEPED electron detectors:
 - **Contamination by energetic protons**
 - **Significantly non-ideal detector response**
- **These problems lead to significant errors and long-term inhomogeneities in the electron data, and require correction.**

Schematic electron
instrument design:

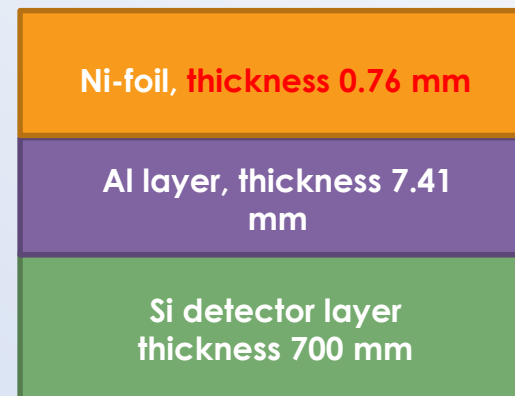


1D detector models (note that
the layers are not in scale):

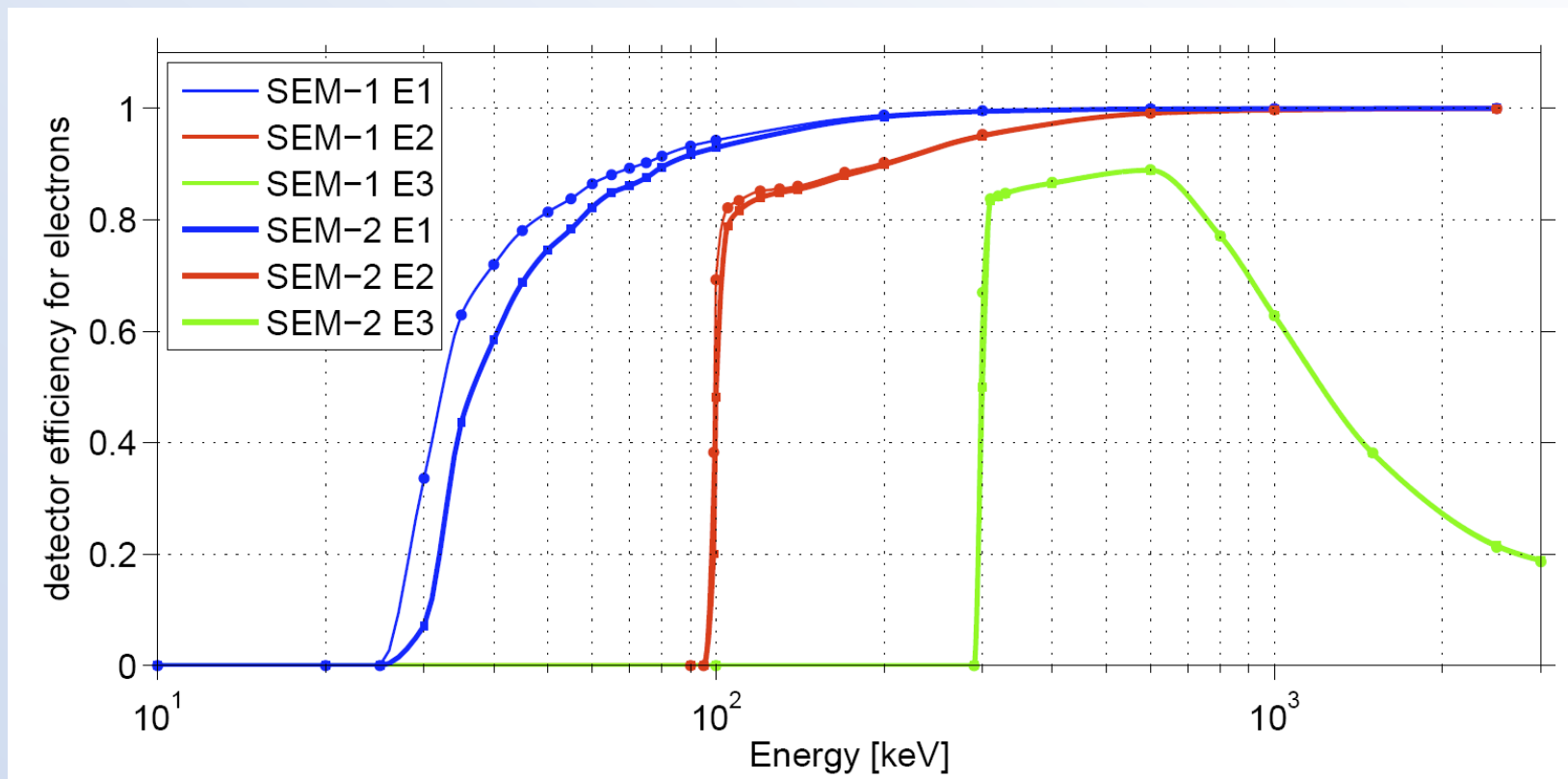
SEM-1 model



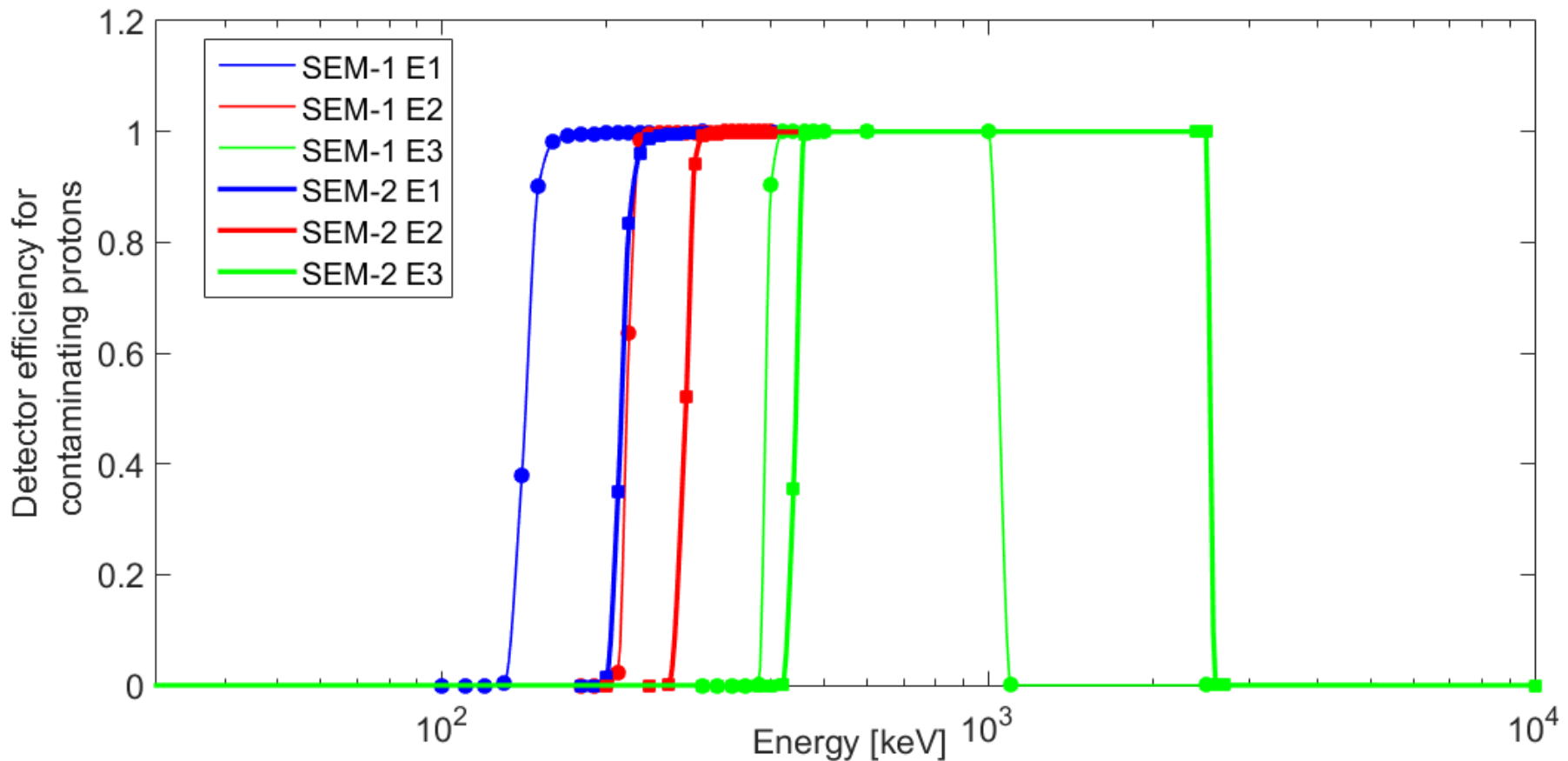
SEM-2 model (note the thicker
shielding layers)



- Efficiencies for all channels deviate from ideal
- Large differences esp. in E1 channel between SEM-1 and SEM-2 instruments.
- **The differences in efficiencies result from the thicker Ni-foil in SEM-2**



- Efficiencies for all channels deviate from ideal
- Large differences in all channels
- Differences again due to different shielding thicknesses.



- We assume that the differential electron spectrum is **piecewise continuous power law**:

$$f_e(E) = AE_{xo}^{\gamma_1 - \gamma_2} E^{-\gamma_1}, E < E_{xo}$$
$$f_e(E) = AE^{-\gamma_2}, E \geq E_{xo}$$

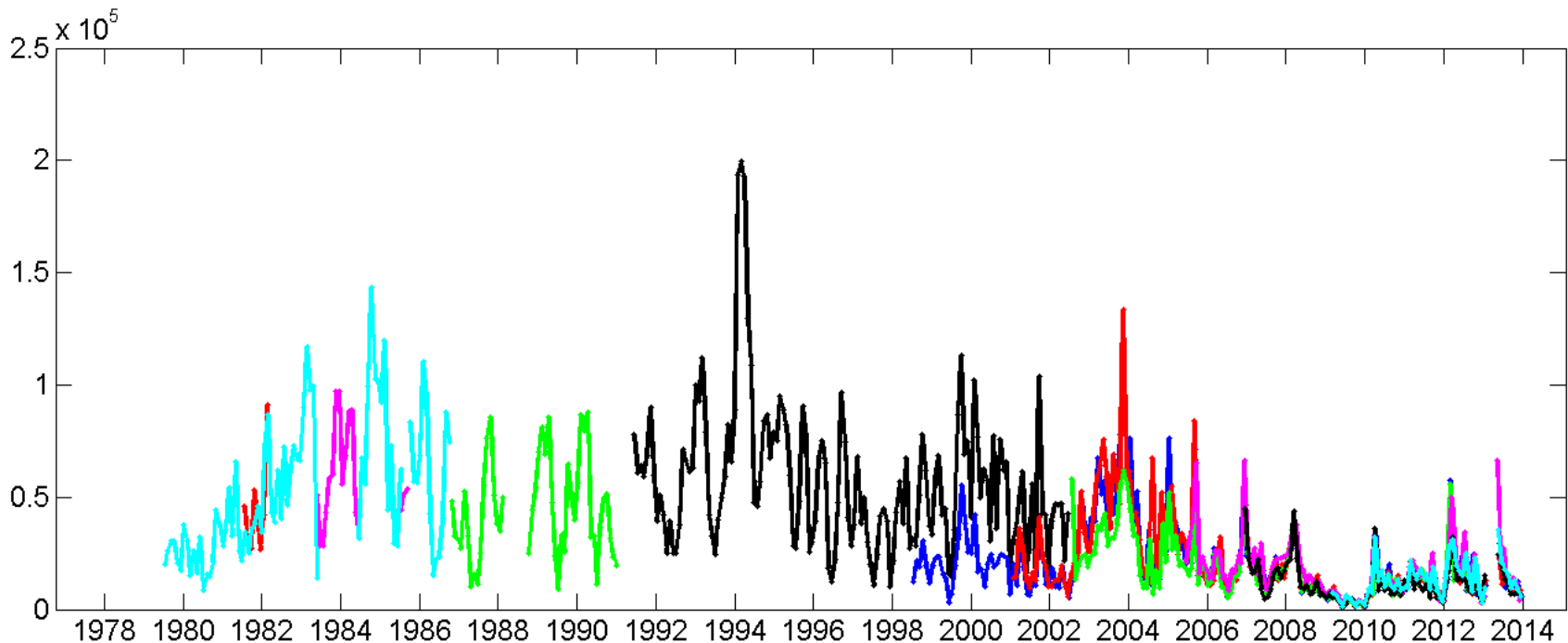
where $E_{xo} = 95$ keV.

- In discrete form the measured fluxes in i :th channel ($i=1,2,3$) can be written as

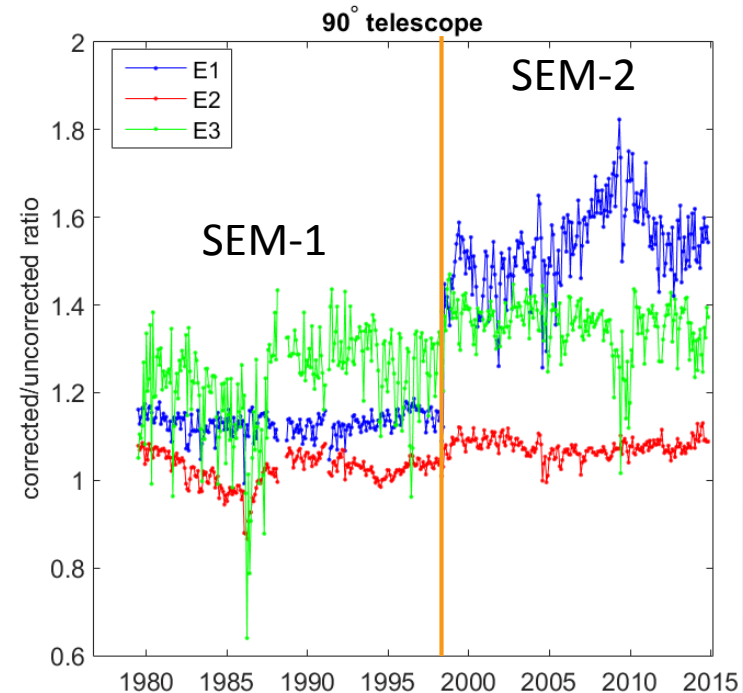
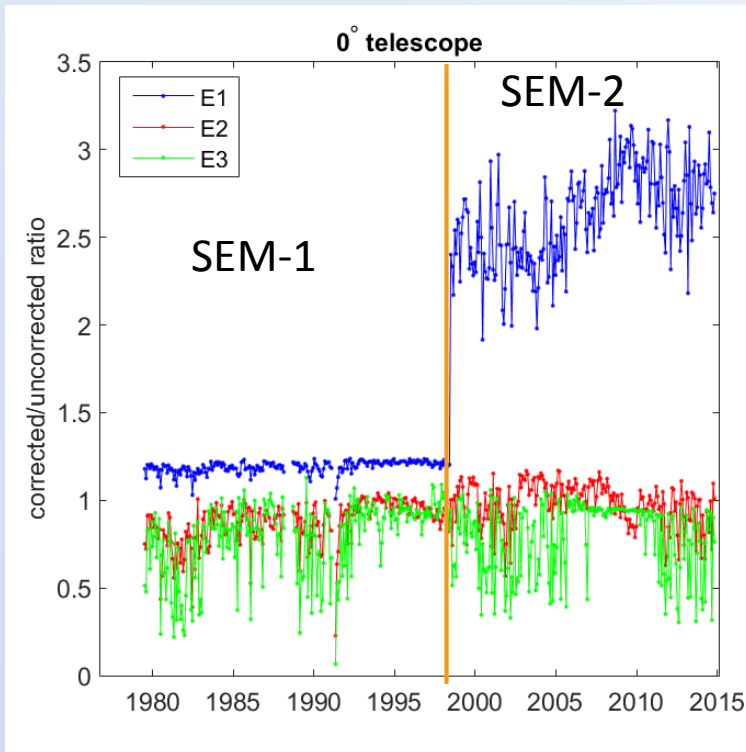
$$j_i = \sum_{E_k=0}^{E_{\max}} \varepsilon_i(E_k) f_e(E_k) \Delta E + \sum_{E_k=0}^{E_{\max}} \rho_i(E_k) f_p(E_k) \Delta E$$

- where $\varepsilon_i(E)$ and $\rho_i(E)$ are the electron and proton efficiency functions for the i :th channel, $E_{\max} = 2.5$ MeV, $\Delta E = 1$ keV, and the differential proton spectrum $f_p(E)$ is obtained from corrected proton measurements.
- **→ Numerically solve for A , γ_1 and γ_2 and use them to compute the fluxes at nominal channel energies (>30 keV, >100 keV and >300 keV)** by integrating the differential spectrum.

- Newer SEM-2 satellites (since mid-1998) systematically smaller than SEM-1 satellites

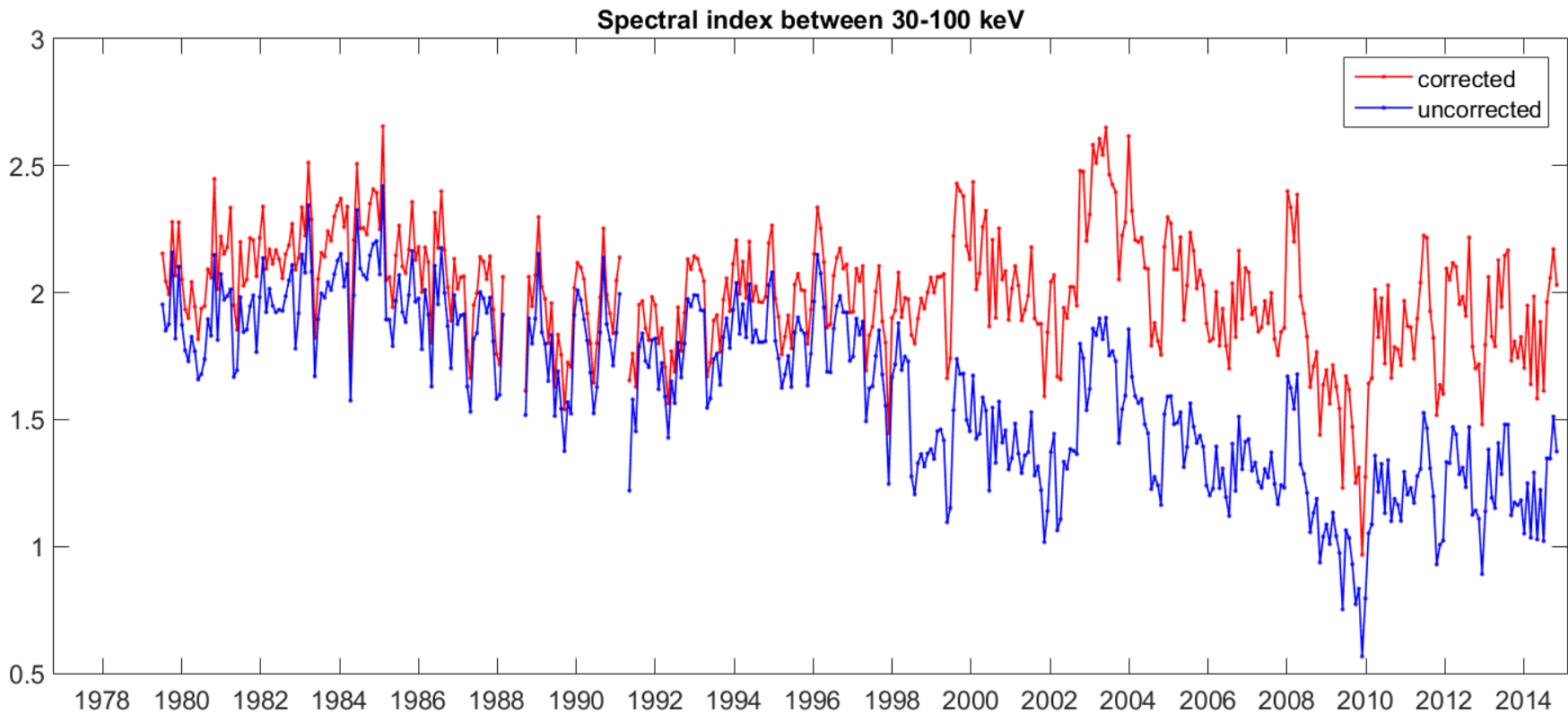


- Correction has two parts: removing proton contamination and correcting for detector sensitivity.
- **→ Overall effect of correction is different in different energy channels, telescopes and instrument versions (SEM-1/SEM-2)**



Energy spectrum is severely distorted without correction

- Erroneous long-term evolution of spectral index
- Energy spectrum too hard without correction



- Data from different sources is typically always inhomogeneous
- Constant difference or temporally changing?
- → Understand your data!!

- Calibrate by comparing simultaneous and comparable measurements
- More refined calibrations and homogenization require deep understanding of instruments and how their properties change with time
- Last but not least: Estimate the uncertainty!

THE END